# GPS, Geodesy, and the Ghost in the Machine 

A Workshop for Surveyors and GIS Professionals

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Workbook state edition: Arizona, version 8
March 2010

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## WORKSHOP ABSTRACT

Coordinates derived from GPS equipment are determined using complex algorithms that are often hidden within proprietary software - the "ghost in the machine". Users unfamiliar with the computational process can unwittingly generate positional errors ranging from a fraction of a foot to many miles. This problem persists despite efforts of vendors to streamline and simplify the GPS positioning process.

This workshop seeks to shed light on the GPS "black box" by 1) Explaining the main geodetic principles and terminology behind GPS; 2) Reducing blind reliance on GPS and GIS software; and 3) Providing practical information and tools for the GPS user. Topics include geodetic and vertical datums, map projections, "ground" coordinate systems, the geoid, NGS Datasheets and OPUS, GIS data compatibility, and an overview of (draft) APLS standards for spatial data accuracy and georeferencing. Numerous examples of positioning errors will be used to illustrate the peril of neglecting geodetic principles in modern surveying and mapping applications. A workbook will be provided that includes step-by-step GPS and geodetic computations. So bring your questions and your favorite everyday tools (calculator, laptop computer, data collector), and together we will purge the ghosts from your machines!

## ACKNOWLEDGEMENTS

The creation of this workbook would not have been possible without the support and inspiration provided by Gabriel Bey and Rick Bunger. Special thanks also go to Dave Minkel, the National Geodetic Survey Arizona Geodetic Advisor, and to all members of the Geospatial Committee of the Arizona Professional Land Surveyors Association.

Today, GPS has thrust surveyors into the thick of geodesy which is no longer the exclusive realm of distant experts. Thankfully, in the age of microcomputers, the computational drudgery can be handled with software packages. Nevertheless, it is unwise to venture into GPS believing that knowledge of the basics of geodesy is, therefore, unnecessary. It is true that GPS would be impossible without computers, but blind reliance on the data they generate eventually leads to disaster.

Jan Van Sickle (2001, p. 126)

Note: This workbook is intended to accompany a presentation. Therefore some of the material may appear incomplete or be unclear if it is used without attending the presentation.

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## LIST OF SYMBOLS

a Semi-major axis of geodetic ellipsoid; also semi-major axis of error ellipse
A Astronomic azimuth
$b \quad$ Semi-minor axis of geodetic ellipsoid, $b=a(1-f)$; also semi-minor axis of error ellipse
$c_{X}^{n} \quad$ Value used to scale an $n$-dimensional standard error to a confidence level of $X \%$
$C$ Geopotential number
$C_{C} \quad$ Horizontal curvature correction factor (multiplied with straight horizontal distance)
$C_{C R} \quad$ Vertical correction for curvature and refraction (added to vertical distance)

## CEP Circular Error Probable

$D C_{A B}$ Dynamic correction applied to leveled height difference from point $A$ to $B$
$D_{\text {grnd }}$ Horizontal ground distance (parallel to ellipsoid)
$D_{S} \quad$ Slope distance
$e \quad$ First eccentricity of geodetic ellipsoid, $e=\sqrt{e^{2}}=\sqrt{2 f-f^{2}}$
$e^{2} \quad$ First eccentricity squared of geodetic ellipsoid, $e^{2}=2 f-f^{2}$
$E \quad$ Easting coordinate (in the $x$-direction)
$E_{95}^{N} \quad$ Error (accuracy) north component at accuracy at $95 \%$ confidence
$E_{95}^{E} \quad$ Error (accuracy) east component at accuracy at $95 \%$ confidence
$E_{95}^{U} \quad$ Error (accuracy) up component at accuracy at $95 \%$ confidence
$E_{0} \quad$ False easting (on central meridian) of map projection definition
$f \quad$ Geometric flattening of geodetic ellipsoid
$g \quad$ Gravity at the topographic surface of the Earth
$\bar{g} \quad$ Mean gravity on the plumbline between the topographic surface and the geoid
$h \quad$ Ellipsoid height
$H \quad$ Orthometric height ("elevation")
$H^{D} \quad$ Dynamic height
$i \quad$ Instrument or GPS base station antenna height above station
$K \quad$ Constant for computing Helmert mean gravity, $K=2,358,000 \mathrm{~s}^{2}=1 /\left(4.24 \times 10^{-7} \mathrm{~s}^{-2}\right)=$
$1 /(1 / 2 F-2 \pi G \rho)$, where $F$ is the vertical gradient of gravity, $G$ is the universal gravitational constant and $\rho$ is the topographic density (assumed constant $2670 \mathrm{~kg} / \mathrm{m}^{3}$ )
$k \quad$ Conformal map projection grid scale factor
$k_{0} \quad$ Grid scale factor on the central meridian for the Transverse Mercator projection (or on the central parallel for the Lambert Conformal Conic projection)
$L \quad$ Laplace correction
$N \quad$ Northing coordinate (in the $y$-direction)
$N_{0} \quad$ False northing (where central meridian crosses latitude of grid origin) of map projection
$N_{G} \quad$ Geoid height $=$ geoid separation $=$ geoid undulation
$\Delta n_{A B}$ Leveled height difference observed from point $A$ to $B$
$O C_{A B}$ Orthometric correction applied to leveled height difference from point $A$ to $B$
$r \quad$ Prism rod or rover antenna height above station
$R_{G} \quad$ Geometric mean radius of curvature of geodetic ellipsoid
$R_{M} \quad$ Radius of curvature in the meridian of geodetic ellipsoid
$R_{N} \quad$ Radius of curvature in the prime vertical of geodetic ellipsoid
$R_{\alpha} \quad$ Radius of curvature of geodetic ellipsoid in a specific azimuth, $\alpha$
$s \quad$ Geodesic distance ("horizontal" distance on the ellipsoid)
SEP Spherical Error Probable
$t$ Grid azimuth
$(t-T)$ Arc-to-chord ("second term") correction for converting grid to geodetic azimuths
$X \quad$ Earth-Centered, Earth-Fixed Cartesian coordinate in the $X$-direction (in equatorial plane and passing through Prime Meridian, i.e., $0^{\circ}$ longitude)
$Y \quad$ Earth-Centered, Earth-Fixed Cartesian coordinate in the $Y$-direction (in equatorial plane and perpendicular to $X$-axis, i.e., passing through $90^{\circ} \mathrm{E}$ longitude)
$Z \quad$ Earth-Centered, Earth-Fixed Cartesian coordinate in the Z-direction (parallel to Earth's conventional spin axis and perpendicular to equatorial plane)
$\alpha_{A B} \quad$ Geodetic azimuth from point $A$ to point $B$
$\widetilde{\alpha}_{A B} \quad$ Approximate geodetic azimuth from point $A$ to point $B$
$\gamma \quad$ Convergence angle
$\gamma_{0} \quad$ Normal gravity on the GRS 80 ellipsoid at $45^{\circ}$ latitude
$\delta \quad$ Map projection distortion
$\Delta \quad$ Denotes discrete change in a quantity, usually as final value minus initial value (e.g., for change in northing coordinate, $\Delta N=N_{2}-N_{1}$ )
$\Delta \lambda^{\prime \prime} \quad$ Change in longitude in arc-seconds
$\Delta \varphi^{\prime \prime} \quad$ Change in latitude in arc-seconds
$\zeta \quad$ Geodetic zenith angle
$\eta \quad$ East-west component of the deflection of the vertical (in the prime vertical plane)
$\theta \quad$ Horizontal error ellipse rotation angle
$\lambda \quad$ Geodetic longitude
$\lambda_{0} \quad$ Longitude of central meridian for map projection
$v$ Zenith angle
$\xi \quad$ North-south component of the deflection of the vertical (in the meridian plane)
$\pi \quad$ Irrational number pi (ratio of circle circumference to diameter)
$\rho \quad$ Horizontal correlation
$\sigma_{E} \quad$ Standard error (east component)
$\sigma_{N} \quad$ Standard error (north component)
$\sigma_{N E}$ Horizontal covariance
$\varphi \quad$ Geodetic latitude (on ellipsoid or sphere)
$\varphi_{0} \quad$ Latitude of grid origin for map projection; central parallel for conical map projection
$\varphi_{\mathrm{N}} \quad$ Latitude of north standard parallel for conical map projection
$\varphi_{\mathrm{S}} \quad$ Latitude of south standard parallel for conical map projection
$\psi \quad$ Angle between two points on a sphere with vertex at center of sphere

TABLE OF USEFUL NUMERICAL VALUES

| Symbol | Description | Numerical values |
| :---: | :---: | :---: |
| $a$ | GRS-80 ellipsoid semi-major axis (identical to WGS-84 value) | $\begin{aligned} 6,378,137 \mathrm{~m}(\text { exact }) & =20,925,646.325459 \mathrm{ift} \\ & =20,925,604.474167 \mathrm{ftt} \end{aligned}$ |
| $f$ | GRS-80 geometrical flattening WGS-84 geometrical flattening | $298.257222101^{-1}$ (published value) $298.257223563^{-1}$ (published value) |
| $b$ | GRS-80 ellipsoid semi-minor axis <br> WGS-84 ellipsoid semi-minor axis | $\begin{aligned} 6,356,752.314140 \mathrm{~m} & =20,855,486.594949 \mathrm{ift} \\ & =20,855,444.883876 \mathrm{sft} \\ 6,356,752.314245 \mathrm{~m} & =20,855,486.595293 \mathrm{ift} \\ & =20,855,444.884319 \mathrm{sft} \end{aligned}$ |
| $e^{2}$ | GRS-80 first eccentricity squared WGS-84 first eccentricity squared | $\begin{aligned} & 0.006694380022901 \\ & 0.006694379990141 \end{aligned}$ |
| ift | International Foot | $1 \mathrm{ift} \equiv 0.3048 \mathrm{~m}(2 \mathrm{ppm}$ shorter than sft$)$ |
| sft | US Survey Foot | $1 \mathrm{sft} \equiv 1200 / 3937 \mathrm{~m}(2 \mathrm{ppm}$ longer than ift) |
| ppm | Parts per million | Value multiplied by one million (analogous to "percent" which is "parts per hundred") |
| rad | Radian (angular measure) | $1 \mathrm{rad}=180^{\circ} / \pi$ (i.e., $1 \mathrm{rad} \approx 57.295779513^{\circ}$ ) |
| $\pi$ | Pi (irrational number) | $\pi=3.141592653589793238462643383 \ldots$ |
| $\gamma_{0}$ | Normal gravity on the GRS 80 ellipsoid at $45^{\circ}$ latitude | $\begin{gathered} 9.806199 \mathrm{~m} / \mathrm{s}^{2} \\ 32.172569 \mathrm{ift} / \mathrm{s}^{2}=32.172505 \mathrm{sft} / \mathrm{s}^{2} \end{gathered}$ |

## Section 1

## GPS, GEODESY, AND THE PERILS OF MODERN POSITIONING

## Exercise 1.1: Computation of coordinates from total station data

Total stations determine three-dimensional coordinates by measuring three quantities: 1) slope distance, 2) horizontal angle, and 3) zenith angle.
Grid coordinates (northing and easting) and elevation can be computed from a total station using the following formulas (designated as Equation 1.1):

Equation 1.1 Computation of grid coordinates from total station data

$$
\begin{aligned}
& N=N_{0}+D_{S} \cos \alpha \sin v \\
& E=E_{0}+D_{S} \sin \alpha \sin v \\
& H=H_{0}+D_{S} \cos v+i-r
\end{aligned}
$$

where $N, E$, and $H$ are the northing, easting, and height (elevation) coordinates to be determined $N_{0}, E_{0}$, and $H_{0}$ are the northing, easting, and height of the instrument setup point $D_{S}$ is the observed slope distance
$\alpha$ is the observed horizontal angle (azimuth)
$v$ is the observed zenith angle
$i$ and $r$ are the instrument and the prism rod heights, respectively.


## Example computation

Given: A total station set up with $i=5.32 \mathrm{ft}$ over starting point with $N_{0}=5000.00 \mathrm{ft}, E_{0}=$ 5000.00 ft , and $H_{0}=100.00 \mathrm{ft}$. The horizontal circle is set so that it reads azimuth directly, and the following observations are made to a point with prism rod of height $r=6.56 \mathrm{ft}$ :

$$
D_{S}=336.84 \mathrm{ft} \quad \alpha=152^{\circ} 17^{\prime} 23^{\prime \prime} \quad v=83^{\circ} 48^{\prime} 50^{\prime \prime}
$$

Find: The coordinates and elevation of the observed point.
Computations:

| $N=$ | $N_{0}$ | $+$ | $D_{S}$ | $\times$ | $\cos \alpha$ | $\times$ | $\sin v$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $N=$ |  | + |  | $\times \cos ($ |  | $) \times \sin ($ |  | _) |
| $N=$ |  | + |  | $\times$ |  | $\times$ |  |  |
| $N=$ |  |  |  |  |  |  | $\sin v$ |  |
| $E=$ | $E_{0}$ | $+$ | $D_{S}$ | $\times$ | $\sin \alpha$ | $\times$ |  |  |
| $E=$ |  | ++ |  | $\times \sin ($ |  | $) \times \sin ($ |  | _) |
| $E=$ |  |  |  | $\times$ |  | $\times$ |  |  |
| $\underline{E}=$ |  |  |  |  |  |  |  |  |
| $H=$ | $H_{0}$ | + | $D_{S}$ | $\times$ | $\cos v$ | $+\quad i$ | - | $r$ |
| $H=$ |  | $+$ |  | $\times \cos ($ |  | ) + | -- |  |
| $H=$ |  | $+$ |  | $\times$ |  | $+$ | - |  |
| $\underline{H}=$ |  |  |  |  |  |  |  |  |

## Solution:

$$
\begin{aligned}
& N=5000.00+336.84 \times \cos \left(152^{\circ} 17^{\prime} 23^{\prime \prime}\right) \times \sin \left(83^{\circ} 48^{\prime} 50^{\prime \prime}\right) \\
& N=5000.00+336.84 \times(-0.88531023) \times 0.99417711 \\
& \underline{\boldsymbol{N}=\mathbf{4 7 0 3 . 5 3} \mathbf{f t}}
\end{aligned}
$$

$$
\begin{aligned}
& E=5000.00+336.84 \times \sin \left(152^{\circ} 17^{\prime} 23^{\prime \prime}\right) \times \sin \left(83^{\circ} 48^{\prime} 50^{\prime \prime}\right) \\
& E=5000.00+336.84 \times 0.4650009 \times 0.99417711 \\
& \underline{\boldsymbol{E}}=\mathbf{5 1 5 5 . 7 2} \mathbf{~ f t}
\end{aligned}
$$

$$
H=100.00+336.84 \times \cos \left(83^{\circ} 48^{\prime} 50^{\prime \prime}\right)+5.32-6.56
$$

$$
H=100.00+336.84 \times 0.10775836+5.32-6.56
$$

$$
H=135.06 \mathrm{ft}
$$

## GPS: A geodetic tool

## A comparison between total stations and GPS

Both GPS and total stations determine three-dimensional coordinates, but they differ in virtually every other respect, to wit:

- Observations
- Total stations are used to directly observe slope distance, horizontal angle, and zenith angle
- Total station EDM sends and receives the signal that it uses for computing distance
- GPS observes the pseudorange, carrier phase (fractional wavelength), and Doppler shift of the signals transmitted from the satellites
- GPS only receives signals from the satellites (a one-way ranging system)
- Measurements
- The vector components from a total station to the prism are directly measured
- Total station measures both distance and angles
- The vector components between GPS antennas are computed, NOT observed
- This has implications for error propagation and control network design
- GPS does NOT measure angles
- Computations
- Coordinates can be determined from total station observations using simple plane trigonometry
- Geodetic methods MUST be used to compute coordinates from GPS vectors
- Reference frame
- Total stations are referenced to the gravity vector (plumbline) passing through the vertical axis of the instrument
- GPS is referenced to a world-wide coordinate system (in common with the satellites) with its origin located at the Earth's center of the mass


## Geodesy: The science of positioning

Geodesy is a quantitative scientific field dealing with the size and shape of the Earth (or other planetary bodies), precise determination of coordinates and relationship between coordinates on the Earth, and includes study of the Earth's gravity field. It is the science behind surveying, mapping, and navigation, and it is essential for using GPS.

The bottom line: GPS is a geodetic tool that requires geodesy to perform computations and it is explicitly referenced to the entire Earth.

The geodetic ellipsoid of revolution


## Earth-Centered, Earth-Fixed (ECEF) Cartesian coordinates

## Ellipsoid



## Exercise 1.2: Geodetic ellipsoid parameters and computations

The geodetic ellipsoid of revolution is completely defined by two numbers. By convention, these are usually $a$, the semi-major axis, and $1 / f$, the inverse geometric flattening. These can be used to compute other commonly used ellipsoid parameters, such as the following two:

Equation 1.2 Ellipsoid semi-minor axis Equation 1.3 Ellipsoid first eccentricity squared

$$
b=a(1-f)
$$

$$
e^{2}=2 f-f^{2}
$$

## Example computations

Given: The following parameters for the GRS-80, WGS-84, and Clarke 1866 ellipsoids:

| Ellipsoid | GRS-80 | WGS-84 | Clarke 1866 |
| ---: | :---: | :---: | :---: |
| Semi-major axis, $\boldsymbol{a}$ | $6,378,137 \mathrm{~m}$ (exact) | $6,378,137 \mathrm{~m}$ (exact) | $20,925,832.164 \mathrm{sft}$ |
| Inverse flattening, $\mathbf{1} \boldsymbol{f}$ | 298.257222101 | 298.257223563 | 294.978698214 |

Find: The semi-minor axis (in International Feet) of these ellipsoids.
Computations:
Semi-minor axis $=a \quad \times \quad(1-f) \quad \times$ unit conversion

$b=\ldots$ ift
WGS-84: $\quad b=\square \times\left[1-\left(\frac{1}{}\right)\right] \times\left(\frac{1 \mathrm{ift}}{0.3048 \mathrm{~m}}\right)$ $b=\longrightarrow$ ift

Clarke 1866: $b=\square \times\left[1-\left(\frac{1}{1}\right)\right] \times\left(\frac{1.000002 \mathrm{ift}}{1 \mathrm{sft}}\right)$

$$
b=
$$

$\qquad$
Solution:
GRS-80: $\quad b=6,378,137 \mathrm{~m} \times\left[1-\frac{1}{298.257222101}\right] \times\left(\frac{1 \mathrm{ift}}{0.3048 \mathrm{~m}}\right)=\underline{\mathbf{2 0 , 8 5 5}, \mathbf{4 8 6 . 5 9 4 9} \mathbf{i f t}}$
WGS-84: $b=6,378,137 \mathrm{~m} \times\left[1-\frac{1}{298.257223563}\right] \times\left(\frac{1 \mathrm{ift}}{0.3048 \mathrm{~m}}\right)=\underline{\mathbf{2 0 , 8 5 5}, \mathbf{4 8 6 . 5 9 5 3} \mathbf{~ i f t}}$
Clarke 1866: $b=20,925,832.164 \mathrm{sft} \times\left[1-\frac{1}{294.978698214}\right] \times \frac{1.000002 \mathrm{ift}}{1 \mathrm{sft}}=\underline{\mathbf{2 0 , 8 5 4 , 9 3 3 . 7 2 7} \mathbf{~ i f t}}$

## Exercise 1.3: Computation of Earth radius

The Radii of curvature at a point in the meridian (north-south) and prime vertical (east-west) are frequently used in geodesy:

Equation 1.4 Meridian radius (north-south)

$$
R_{M}=\frac{a\left(1-e^{2}\right)}{\left(1-e^{2} \sin ^{2} \varphi\right)^{3 / 2}}
$$

Equation 1.5 Prime vertical radius (east-west)

$$
R_{N}=\frac{a}{\sqrt{1-e^{2} \sin ^{2} \varphi}}
$$

where $\varphi$ is the geodetic latitude at the point where the radius is computed.
$a$ is the ellipsoid semi-major axis $(=20,925,646.325459 \mathrm{ift}$ for the GRS-80 ellipsoid)
$e^{2}$ is the ellipsoid first eccentricity squared $(=0.006694380022901$ for GRS-80)
$R_{M}$ and $R_{N}$ are used to compute other commonly used Earth radii, such as the following two:
Equation 1.6 Radius of curvature in a specific azimuth, $\alpha$

$$
R_{\alpha}=\frac{R_{M} R_{N}}{R_{M} \sin ^{2} \alpha+R_{N} \cos ^{2} \alpha}
$$

Equation 1.7 Geometric mean radius of curvature

$$
R_{G}=\sqrt{R_{M} R_{N}}=\frac{a \sqrt{1-e^{2}}}{1-e^{2} \sin ^{2} \varphi}
$$

$R_{G}$ is the essentially the "average" radius of curvature at a point on the ellipsoid, and is the one we will use for radius computations in this workshop.

Some $\boldsymbol{R}_{\boldsymbol{G}}$ values for Arizona:

$B$
Rule of thumb: Geometric mean radius of curvature increases by about 33 ft per mile north (between latitudes of $30^{\circ}$ and $40^{\circ}$ )

## Example computation

Given: A point at latitude $\varphi=34^{\circ} 32^{\prime} 59.29087^{\prime \prime} \mathrm{N}$ (midway between points CAS-2 and CAS-3).
Find: The radii of curvature in the meridian, prime vertical, at an (approximate) azimuth of $\alpha=$ $72^{\circ} 06^{\prime} 17^{\prime \prime}$ (from CAS-2 to CAS-3), and the geometric mean radius (for the GRS-80 ellipsoid).
Computations: First convert latitude and azimuth to decimal degrees:

$$
\begin{aligned}
& \varphi=34+32 / 60+59.29087 / 3600=34.5498030194^{\circ} \\
& \alpha=72+6 / 60+17 / 3600=72.10473^{\circ}
\end{aligned}
$$

Now compute following function of latitude (since it appears in most of the equations):

$$
1-e^{2} \sin ^{2} \varphi=1-0.006694380023 \times\left[\sin \left(34.5498030194^{\circ}\right)\right]^{2}=0.99784690136
$$

Now compute the various radii:

$\qquad$

$\qquad$

$$
R_{G}=\overline{\times \sqrt{1-}}=
$$

$\qquad$

Solution:

$$
\begin{array}{ll}
R_{M}=\frac{20,925,646.325 \times(1-0.006694380023)}{(0.99784690136)^{3 / 2}} & =\underline{\mathbf{2 0 , 8 5 2}}, \\
R_{N}=\frac{20,925,646.325}{\sqrt{0.99784690136}} & =\underline{\mathbf{2 0 , 9 4 8}}, \\
R_{\alpha}=\frac{20,852,872.616 \times 20,948,210.040}{20,852,872.616 \times\left[\sin \left(72.10473^{\circ}\right)\right]^{2}+20,948,210.040 \times\left[\cos \left(72.10473^{\circ}\right)\right]^{2}}
\end{array}
$$

$$
=\underline{20,939,171.046} \text { ift }
$$

$$
R_{G}=\frac{20,925,646.325 \times \sqrt{1-0.006694380023}}{0.99784690136} \quad=\underline{\mathbf{2 0 , 9 0 0}, \mathbf{4 8 7 . 4 0 6} \mathbf{~ i f t}}
$$

Check: $R_{G}=\sqrt{R_{M} R_{N}}=\sqrt{20,852,873 \times 20,948,210}=\underline{20,900,487.406 \mathrm{ift}} \downarrow$

## The NGS Datasheet



Section 1: GPS, Geodesy, and the Perils of Modern Positioning

The NGS Datasheet (continued)

```
ES0478 SUPERSEDED SURVEY CONTROL
ES0478
ES0478 NAD 83(1992)- 34 43 41.84276(N) 111 58 50.37059(W) AD( ) A
ES0478 ELLIP H (09/30/99) 1000.753 (m) GP( ) 3 1
ES0478 NAVD 88 (09/30/99) 1026.38 (m) 3367.4 (f) LEVELING 
ES0478 NGVD 29 (??/??/92) 1025.631 (m) 3364.92 (f) ADJ UNCH 1 2
ES0478
ES0478.Superseded values are not recommended for survey control.
ES0478.NGS no longer adjusts projects to the NAD 27 or NGVD 29 datums.
ES0478. See file dsdata.txt to determine how the superseded data were derived.
ES0478
ES0478_U.S. NATIONAL GRID SPATIAL ADDRESS: 12SVD1021643349(NAD 83)
ES0478_MARKER: DD = SURVEY DISK
ESO478_SETTING: 66 = SET IN ROCK OUTCROP
ES0478_SP_SET: LIMESTONE LEDGE
ES0478_STAMPING: R 18-1931
ES0478 MARK LOGO: USGS-E
ESO478-MAGNETIC: N = NO MAGNETIC MATERIAL
ES0478_STABILITY: A = MOST RELIABLE AND EXPECTED TO HOLD
ESO478+STABILITY: POSITION/ELEVATION WELL
ESO478_SATELLITE: THE SITE LOCATION WAS REPORTED AS SUITABLE FOR
ES0478+SATELLITE: SATELLITE OBSERVATIONS - April 25, 2009
ES0478
ES0478 HISTORY - Date Condition Report By
ES0478 HISTORY - UNK MONUMENTED USGS-E
ES0478 HISTORY - 1933 GOOD NGS
ES0478 HISTORY - 19990121 GOOD AZ-025
ES0478 HISTORY - 20031215 GOOD SHEPH
ES0478 HISTORY - 20090425 GOOD GEOCAC
ES0478
ES0478 STATION DESCRIPTION
ES0478
ESO478'DESCRIBED BY NATIONAL GEODETIC SURVEY 1933
ESO478'4.4 MI SE FROM COTTONWOOD.
ESO478'AT SIDE OF HIGHWAY, 0.8 MILES NORTH OF VERDE RIVER HIGHWAY BRIDGE, TOP
ESO478'OF RIDGE, 100 FEET NORTHWEST OF HIGHWAY CENTER-LINE, AT SIDE OF ROCK
ES0478'CAIRN, ON LEDGE OF LIMESTONE PAINTED BLACK U.S.B.M. 3363.9.
ES0478
ES0478 STATION RECOVERY (1999)
ES0478
ESO478'RECOVERY NOTE BY YAVAPAI COUNTY ARIZONA 1999 (WRA)
ES0478'THE STATION IS LOCATED ABOUT 4.4 MI (7.1 KM) SOUTHEAST OF COTTONWOOD,
ES0478'1 MI (1.6 KM) NORTHEAST OF BRIDGEPORT, 0.8 MI (1.3 KM) NORTHEAST OF A
ES0478'HIGHWAY BRIDGE OVER THE VERDE RIVER, 0.35 MI (0.56 KM) SOUTHWEST OF
ES0478'CORNVILLE ROAD, 0.25 MI (0.40 KM) NORTHEAST OF ROCKING CHAIR ROAD, AT
ES0478'US HIGHWAY 89 ALTERNATE MILEPOST 356.7. OWNERSHIP--COCONINO NATIONAL
ESO478'FOREST. TO REACH THE STATION FROM THE JUNCTION OF U.S. HIGHWAY 89
ESO478'ALTERNATE AND STATE HIGHWAY 279 IN COTTONWOOD, GO NORTHEAST FOR 2.2 KM
ES0478'(1.35 MI) ON HIGHWAY 89 ALTERNATE TO THE STATION ON RIGHT, AT THE TOP
ESO478'OF A SMALL KNOLL. THE STATION IS A DISK SET IN A LIMESTONE OUTCROP.
ES0478'LOCATED 41.5 M (136.2 FT) SOUTHEAST FROM THE CENTERLINE OF HIGHWAY,
ES0478'32.0 M (105.0 FT) NORTHWEST FROM THE CENTER OF AN ABANDONED ROAD, 21.7
ES0478'M (71.2 FT) SOUTHEAST FROM A FENCE, 0.9 M (3.0 FT) SOUTH-SOUTHEAST
ESO478'FROM A ROCK CAIRN AND 0.3 M (1.0 FT) WEST FROM A WITNESS POST.
ES0478
ES0478 STATION RECOVERY (2003)
ES0478
ES0478'RECOVERY NOTE BY SHEPHARD-WESNITZER INC 2003 (MLD)
ES0478'RECOVERED AS DESCRIBED
ES0478
ES0478 STATION RECOVERY (2009)
ES0478
ESO478'RECOVERY NOTE BY GEOCACHING 2009 (ACM)
ESO478'RECOVERED IN GOOD CONDITION.
```


## The NGS Geodetic Toolkit


on-line interactive computation of geodetic values

See the text version of an article about the NGS Geodetic Toolkit that appeared in the Professional Surveyor magazine, May 2003 Volume 23, Number 4
(See all the Professional Surveyor Articles about the NGS Geodetic Toolkit )
To learn more about a particular online program, click on its link for a description:

| DEFLEC99 | Inverse/Forward/Invers3D/Forwrd3D | Surface Gravity Prediction |
| :---: | :---: | :---: |
| DEFLEC09 | LVL DH | Tidal and Orthometric Elevations. |
| DYNAMIC_HT | Magnetic Declination | U.S. National Grid |
| GEOID09 | NADCON | Universal Transverse Mercator Coordinates |
| GEOID03 | NAVD 88 Modelled Gravity | VERTCON |
| GEOID99 | Online Adjustment User Services | XYZ Coordinate Conversion |
| G99SSS | Online Adjustment Utilities User Services |  |
| USGG2009 | OPUS |  |
| USGG2003 | State Plane Coordinates |  |
| $\begin{aligned} & \text { HTDP } \\ & \text { IGLD85 } \end{aligned}$ |  |  |

OR... Know what you want to do?
Select a function from this list:
SELECT A TOOLKIT SHORTCUT

For more information contact NGS Information Services: by e-mail, or call (301) 713-3242, Monday - Friday, 7:00 AM - 4:30 PM eastern time.
http://www.ngs.noaa.gov/TOOLS/ Web site owner: National Geodetic Survey Last updated by NGS.Webmaster on Thursday, 18-Feb-2010 16:05:30 EST

## Section 2

## GEODETIC DATUM DEFINITIONS AND REFERENCE COORDINATES

## How are the data connected to the Earth?

## Examples of georeferencing errors for Arizona

Table 2.1 Examples of various positioning error sources and their magnitudes for Arizona due to geodetic datum definition and reference coordinate problems (abbreviations and technical terms are defined in the Glossary).

| Positioning error examples for Arizona | Error magnitudes |
| :---: | :---: |
| Using NAD 27 when NAD 83 required | Varies from $\sim 210$ to 230 feet (horizontal) |
| Using "WGS 84" when NAD 83 required (e.g., by using WAAS corrections or CORS ITRF coordinates) | $\sim 4$ feet (horizontal) $\sim 3$ feet (vertical) |
| Using published three-parameter datum transformation between NAD 27 and "WGS 84" for NAD 83 projects | $\sim 2$ to 16 feet (horizontal) |
| Using NADCON to transform coordinates between NAD 27 and NAD 83 | $\sim 1$ foot (horizontal) |
| Using NADCON to transform coordinates between NAD 83(1986) "original" and NAD 83(1992) "HARN" | $\sim 0.5$ foot (horizontal) |
| Using NAD 83(1986) "original" when NAD 83(1992) "HARN" required | Up to 3.8 feet (horizontal) |
| Using NAD 83(1992) "HARN" when CORS or 1999 Arizona FBN unpublished coordinates required | Up to 0.2 foot (horizontal and vertical) |
| Using NAD 83(1992) "HARN" when NAD 83(NSRS2007) "National Readjustment" coordinates required | Up to 0.5 ft (horizontal) Up to 0.7 ft (vertical) |
| Using published NGS 14-parameter transformation between "WGS 84" and NAD 83 (CORS) but ignoring velocities and reference (zero) time of 1997 | $\sim 0.6 \mathrm{ft}$ (horizontal) <br> for coordinates in year 2007 |
| Using reference coordinates found in the header records of CORS raw GPS data files | Varies from zero to over 100 feet (horizontal and vertical) |
| Autonomous (uncorrected) GPS single-point positioning precision (at 95\% confidence) | $\sim 10$ to 20 ft (horizontal) $\sim 20$ to 50 ft (vertical) |

## The NGS Datasheet as a geodetic reference coordinate source

Recommend using Datasheets with GPS-derived coordinates, because they give ellipsoid height (as well as ECEF coordinates).


## Some things to note about NGS Datasheets:

- Units of "feet" in NGS Datasheets are presently US Survey Feet, e.g., for above Datasheet:

NAVD $88 \boldsymbol{H}=\mathbf{2 1 6 8 . 4 8 0} \mathbf{m}=\mathbf{7 1 1 4 . 4 2} \mathbf{~ s f t}=\mathbf{7 1 1 4 . 4 4} \mathbf{i f t}$

- Many conventional stations do not have accurate elevations, so cannot be used with geoid model to determine accurate ellipsoid heights
- Conventionally (optically) determined control is almost always less accurate than surveygrade GPS, so using such control for surveys is not advised
- Only GPS stations included in the NSRS2007 readjustment have positional accuracies given as linear "network" values in centimeters (relative "order" system not used)
- Epoch date may not be same as CORS ("Continuously Operating Reference Station")
- NGS station coordinates were determined in 2007 ("NSRS2007")
- 2007.00 epoch date is used for tectonically active states (California, Arizona, Nevada, Oregon, Washington, and Alaska)
- 2002.00 epoch date is used for all other states (consistent with CORS epoch date)


## OPUS output as a geodetic reference coordinate source

## The Online Positioning User Service

This is an excellent alternative to the NGS Datasheets if there are no high-quality GPS-derived NGS control stations locally available.

- More accurate than conventional (optical) control
- Requires logging raw GPS data (observables) at the receiver for at least 2 hours (or as little as 15 minutes using the "Rapid Static" option)
- This can easily be done at a GPS base while performing a survey



## Some things to note about OPUS output:

- Gives both NAD 83 and ITRF 00 coordinates.
- NAD 83 is for epoch 2002.0 (same as NSRS2007 for most of US, except for California and parts of Arizona, Nevada, Oregon, Washington, and Alaska).
- ITRF 00 is for day of observation (e.g., date $2005.4361=$ June 9, 2005 for this example).
- This is NOT same as the current version of WGS-84 (G1150), which was computed for a reference time of 2001.0, so the coordinates will differ by the date difference times the ITRF station velocity (about $0.06 \mathrm{ft} / \mathrm{year}$ to the SW in AZ , so for this case nearly 0.3 ft ).
- However, ITRF 00 and WGS-84 (G1150) can be considered equivalent to within about $1-2 \mathrm{~cm}(0.03-0.07 \mathrm{ft})$ if both refer to the same reference time.
- Slightly different results will be obtained depending on which GPS orbits were used.
- Final orbits available after about 2 weeks.
- "Rapid" orbits available in 17 hours, and are nearly as accurate as final orbits.
- Values to right of coordinates are accuracy estimates in meters, e.g., 0.020 (m).
- These are based on the maximum difference between the 3 positions computed by OPUS.
- Can also estimate accuracy (or at least precision) yourself if have multiple OPUS solutions on a single point.
- Detailed ("extended") output also available
- Gives additional information such as CORS details, coordinate transformations, velocities, actual vector components, GPS solution statistics, and internal precision estimates.
- Two versions of OPUS now available: OPUS-S ("Static") and OPUS-RS ("Rapid Static").
- OPUS-S was formerly simply known as "OPUS" and requires a dataset duration of at least 2 hours.
- OPUS-RS will process shorter datasets (duration from 15 minutes to 2 hours).
- Accuracy of OPUS-RS results varies by location and is best in areas with dense CORS coverage.
- If poor results are achieved with OPUS-RS, use of OPUS-S is recommended (for dataset durations of more than 2 hours).

Relative positioning with "survey-grade" GPS


GPS computation flowchart


## Exercise 2.1: Computation of coordinates using GPS vector components

Below are equations for computing geodetic coordinates of a new station using the GPS vector from a base station of known geodetic coordinates.
Equation 2.1 Converting latitude, longitude, and height to ECEF coordinates

$$
\begin{aligned}
X & =\left(R_{N}+h\right) \cos \varphi \cos \lambda \\
Y & =\left(R_{N}+h\right) \cos \varphi \sin \lambda \\
Z & =\left[R_{N}\left(1-e^{2}\right)+h\right] \sin \varphi
\end{aligned} \text { (Leick, 2004, p. 371) }
$$

where $X, Y$, and $Z$ are the ECEF coordinates of a point
$\varphi, \lambda$, and $h$ are the latitude, longitude, and ellipsoid height of the point, respectively $R_{N}=a\left(1-e^{2} \sin ^{2} \varphi\right)^{-1 / 2}$ is the prime vertical radius of curvature (Leick, 2004, p. 369) $a$ is the ellipsoid semi-major axis ( $=20,925,646.325459$ ift for the GRS-80 ellipsoid) $e^{2}$ is the ellipsoid first eccentricity squared $(=0.006694380022901$ for GRS-80)

Equation 2.2 Computing coordinates from GPS vector components

| $X=X_{b}+\Delta X$ | $Y=Y_{b}+\Delta Y$ | $Z=Z_{b}+\Delta Z$ |
| :--- | :--- | :--- |

where $X, Y$, and $Z$ are the ECEF coordinates to be determined
$X_{b}, Y_{b}$, and $Z_{b}$ are the ECEF coordinates of the GPS base
$\Delta X, \Delta Y$, and $\Delta Z$ are the delta ECEF components of the GPS vector

Equation 2.3 Converting ECEF coordinates to latitude, longitude, and height
$\varphi=\tan ^{-1}\left[\frac{Z}{\sqrt{X^{2}+Y^{2}}}\left(1+\frac{e^{2} R_{N} \sin \varphi_{0}}{Z}\right)\right]$
$\lambda=\tan ^{-1}\left(\frac{Y}{X}\right)$
$h=\frac{\sqrt{X^{2}+Y^{2}}}{\cos \varphi}-R_{N}$$\quad$ (Leick, 2004, pp. 371-372)
where $\varphi_{0}$ is a latitude that can be initially approximated as $\varphi_{0}=\tan ^{-1}\left(\frac{Z}{\left(1-e^{2}\right) \sqrt{X^{2}+Y^{2}}}\right)$.
This approximate latitude value is then substituted into the right side of the first line of Equation 2.3 , and then the resulting value of $\varphi$ is substituted as $\varphi_{0}$, and the process repeated until the change in $\varphi$ is negligible.

## Example computation

Given: A GPS base station located at midpoint between points CAS-2 and CAS-3, with NAD 83 coordinates of $\varphi=34^{\circ} 32^{\prime} 59.29087^{\prime \prime} \mathrm{N}, \lambda=112^{\circ} 26^{\prime} 45.18607^{\prime \prime} \mathrm{W}$, and $h=5456.421 \mathrm{ift}$. The following GPS vector components were determined from this base to point CAS-2:

$$
\Delta X=-219.000 \mathrm{ift} \quad \Delta Y=38.340 \mathrm{ift} \quad \Delta Z=-51.528 \mathrm{ift}
$$

Find: The NAD 83 coordinates of point CAS-2.

## Computations:

Step 1. Convert GPS base latitude, longitude, and ellipsoid height to ECEF coordinates.
The prime vertical radius of curvature for this station was computed in Exercise 1.3:

$$
R_{N}=
$$

$\qquad$
Now compute the ECEF values for the GPS base:

| $X_{b}=($ | $R_{N}$ | + | $h$ | $) \times$ | $\cos \varphi$ | $\times$ | $\cos \lambda$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $X_{b}=($ |  | $+$ |  | ) $\times$ | ( | ) $\times \cos$ |  |
| $Y_{b}=($ | $R_{N}$ | + | $h$ | ) $\times$ | $\cos \varphi$ | $\times$ | $\sin \lambda$ |
| $Y_{b}=($ |  | $+$ |  | _) $\times$ |  | ) $\times \sin ($ |  |
| $Z_{b}=[$ | $R_{N}$ | $\times(1-$ |  | $e^{2}$ | $)+\quad h$ | ] $\times$ | $\sin \varphi$ |
| $Z_{b}=[$ |  | $\times(1-$ |  |  | ) + | ] $\times \sin ($ |  |

Step 2. Compute ECEF coordinates of new GPS station (CAS-2).

$$
\begin{array}{ll}
X=X_{b}+\Delta X= & \\
+ & = \\
Y=Y_{b}+\Delta Y=\square & =\square \\
Z=Z_{b}+\Delta Z=\square & = \\
\hline
\end{array}
$$

Step 3. Convert ECEF coordinates of new station to latitude, longitude, and ellipsoid height.
Instead of using iterative Equation 2.3, perform this computation using the NGS Geodetic Toolkit, which gives the following coordinates for CAS-2:


These results can be verified using Equation 2.3

## Solution:

Step 1. Convert GPS base latitude, longitude, and ellipsoid height to ECEF coordinates.
The prime vertical radius of curvature for this station was computed in Exercise 1.3:

$$
R_{N}=\underline{20,948,210.259 \mathrm{ift}}
$$

Now compute the ECEF values for the GPS base:

$$
\begin{aligned}
& X_{b}=\left(\quad R_{N}+h\right) \times \cos \varphi \times \cos \lambda \\
& X_{b}=(20,948,210.259+5456.421) \times \cos \left(34.5498030194^{\circ}\right) \times \cos \left(-112.4458850194^{\circ}\right) \\
& =\underline{-6,589,343.061 ~ i f t} \\
& Y_{b}=\left(R_{N}+h \quad\right) \times \cos \varphi \times \sin \lambda \\
& Y_{b}=(20,948,210.259+5456.421) \times \cos \left(34.5498030194^{\circ}\right) \times \sin \left(-112.4458850194^{\circ}\right) \\
& =\underline{-15,950,675.460 ~ i f t} \\
& \left.Z_{b}=\left[\begin{array}{llll} 
& R_{N} & \times(1- & e^{2}
\end{array}\right)+h \quad\right] \times \sin \varphi \\
& Z_{b}=[20,948,210.259 \times(1-0.006694380023)+5456.421] \times \sin \left(34.5498030194^{\circ}\right) \\
& =\underline{\mathbf{1 1 , 8 0 3}, 762.654} \mathrm{ift}
\end{aligned}
$$

Step 2. Compute ECEF coordinates of new GPS station (CAS-2).

$$
\begin{aligned}
& X=X_{b}+\Delta X=(-6,589,343.061 \mathrm{ift}) \quad+(-219.000 \mathrm{ift})=\underline{\mathbf{- 6 , 5 8 9}, 562.061 ~ i f t} \\
& Y=Y_{b}+\Delta Y=(-15,950,675.460 \mathrm{ift})+(38.340 \mathrm{ift})=\underline{\mathbf{- 1 5 , 9 5 0 , 6 3 7 . 1 2 0} \mathbf{i f t}} \\
& Z=Z_{b}+\Delta Z=(11,803,762.654 \mathrm{ift}) \quad+(-51.528 \mathrm{ift})=\underline{\mathbf{1 1 , 0 0 3}, 711.126 ~ i f t}
\end{aligned}
$$

Step 3. Convert ECEF coordinates of new station to latitude, longitude, and ellipsoid height.
Equation 2.3 was used to compute the following results for station CAS-2 (compare to those computed using the NGS Geodetic Toolkit).

Latitude, $\varphi=34^{\circ} 32,58.60097 "$ N
Longitude, $\lambda=112^{\circ}$ 26' $47.78016 "$ W
Ellipsoid height, $h=5466.883$ ift

These results were computed using Equation 2.3
(required only 2 iterations in Excel for accuracy shown)

## Datums and datum transformations

Datum. Any quantity or set of quantities used as a reference or basis for determining other quantities.

Geodetic datum. A set of (at least 8) constants specifying the coordinate system for geodetic control (latitude, longitude, height).

2 for reference ellipsoid size and shape (usually semi-major axis and flattening)
3 to specify location of origin (at or near center of Earth for modern datums)
3 to specify the orientation of coordinate system
Vertical datum. A set of fundamental "elevations" to which other "elevations" are referred.
Datum transformation. Mathematical method for converting one geodetic or vertical datum to another (there are several types, and they vary widely in accuracy).

Geodetic datum transformation


Typical geodetic datum transformations. Note that the dimensions of the reference ellipsoid ( $a$ and $b$ axes) may or may not change in the transformation.

3-parameter: 3-dimensional translation of origin as $\Delta X, \Delta Y, \Delta Z$ (like a GPS vector)
7-parameter: 3 translations plus 3 rotations (one about each of the axes) plus a scale
14-parameter: A 7-parameter where each parameter changes with time (each has a velocity)
Transformations are also used that model distortion, such as the NGS model NADCON
Vertical datum transformations. Can be a simple vertical shift or a complex operation that models distortion, such as the NGS model VERTCON.

## Exercise 2.2: Geodetic azimuths

Forward and reverse grid azimuths differ by exactly $180^{\circ}$. Forward and reverse geodetic azimuths do not differ by $180^{\circ}$ because of meridian convergence, as shown in the figure below.

Equation 2.4 Approximate forward geodetic azimuth (from point $A$ to point $B$ )

$$
\tilde{\alpha}_{A B}=\tan ^{-1}\left(\frac{\lambda_{B}-\lambda_{A}}{\varphi_{B}-\varphi_{A}} \cos \varphi_{B}\right)
$$

where $\widetilde{\alpha}_{A B}$ is the approximate forward geodetic azimuth from point $A$ to $B$
$\varphi_{A}, \varphi_{B}$ are longitudes at azimuth end points $A$ and $B$, respectively
$\lambda_{A}, \lambda_{B}$ are longitudes at azimuth end points $A$ and $B$, respectively
Equation 2.4 is accurate to within approximately $\pm 0.5 \%$ for distances of less than about 100 miles.

Although forward and backward grid azimuths differ by exactly $180^{\circ}$, forward and backward geodetic azimuths generally do not due to meridian convergence, as shown in the figure below.


3

## Rule of Thumb:

The average convergence in Arizona is about 35 arc-seconds per mile east-west.
Equation 2.5 Difference between forward and back geodetic azimuths (meridian convergence)

$$
\alpha_{B A}-\alpha_{A B}-180^{\circ} \approx\left(\lambda_{B}-\lambda_{A}\right) \sin \bar{\varphi} \text { (Stem, 1990, p. 51; Ewing and Mitchell, 1970, p. 44) }
$$

where $\alpha_{A B}, \alpha_{B A}$ are the forward and back geodetic azimuths from point $A$ to $B$, respectively $\bar{\varphi}$ is the average latitude of the azimuth end points

Although Equation 2.5 is for a sphere, it is accurate to better than $0.2 "$ for distances of less than about 100 miles.

## Example computation

Given: A two points (CAS-2 and CAS-3) with the following geodetic coordinates:
CAS-2: $\quad \varphi_{A}=34^{\circ} 32^{\prime} 58.60097^{\prime \prime} \mathrm{N} \quad \lambda_{A}=112^{\circ} 26^{\prime} 47.78016^{\prime \prime} \mathrm{W}$
CAS-3: $\quad \varphi_{B}=34^{\circ} 32^{\prime} 59.98077^{\prime \prime} \mathrm{N} \quad \lambda_{B}=112^{\circ} 26^{\prime} 42.59198^{\prime \prime} \mathrm{W}$
Find: The approximate geodetic azimuth from CAS-2 to CAS-3 and compute the difference between the forward and back geodetic azimuths (i.e., the convergence).

## Computations:

To simplify the computations, the approximate geodetic azimuth can be computed using the coordinate differences in arc-seconds:

$$
\begin{aligned}
& \tilde{\alpha}_{A B} \approx \tan ^{-1}\left(\frac{\lambda_{B}-\lambda_{A}}{\varphi_{B}-\varphi_{A}} \cos \varphi_{B}\right)=\tan ^{-1}\left(\frac{(\square)-(\square)}{(\square)} \times \cos ( \right.
\end{aligned}
$$

The difference between forward and back azimuths is

$$
\begin{aligned}
\alpha_{B A}- & \alpha_{A B}-180^{\circ} \approx\left(\lambda_{B}-\lambda_{A}\right) \sin \bar{\varphi} \quad(\text { can use midpoint latitude from Exercises } 1.3 \text { or } 2.1) \\
& =(\square) \times \sin (\square) \\
& =\square
\end{aligned}
$$

Solution:
The approximate geodetic azimuth can be computed as

$$
\begin{aligned}
& \tilde{\alpha}_{A B}=\tan ^{-1}\left(\frac{\lambda_{B}-\lambda_{A}}{\varphi_{B}-\varphi_{A}} \cos \varphi_{B}\right)=\tan ^{-1}\left(\frac{\left(-42.59198^{\prime \prime}\right)-\left(-47.78016^{\prime \prime}\right)}{\left(59.98077^{\prime \prime}\right)-\left(58.60097^{\prime \prime}\right)} \times \cos \left(34.5499946583^{\circ}\right)\right) \\
&=\tan ^{-1}(3.7600957 \times 0.823631645)=\underline{\mathbf{7 2 . 1 0 4 7 3}^{\circ}}=\underline{\mathbf{7 2}^{\circ} \mathbf{0 6}} \mathbf{1 7 ^ { \prime }}=\underline{\mathbf{N}} \mathbf{\mathbf { 7 2 } ^ { \circ } \mathbf { 0 6 }} \mathbf{} \mathbf{1 7} \boldsymbol{1 7} \mathbf{E}
\end{aligned}
$$

The difference between forward and back azimuths is

$$
\begin{aligned}
\alpha_{B A}- & \alpha_{A B}-180^{\circ} \approx\left(\lambda_{B}-\lambda_{A}\right) \sin \bar{\varphi} \quad(\text { can use midpoint latitude from Exercises } 1.3 \text { or } 2.1) \\
& =\left(-42.59198^{\prime \prime}+47.78016^{\prime \prime}\right) \times \sin \left(34.5498030194^{\circ}\right) \\
& =5.18818 " \times 0.5671224=+\mathbf{+ 2 . 9 4 2 3 "}
\end{aligned}
$$

## Check using NGS Inverse tool:

Forward azimuth $=72^{\circ} 10^{\prime} 50.3098^{\prime \prime}$
Error in approximate azimuth is $-0^{\circ} 04^{\prime} 33^{\prime \prime}=-0.11 \%$ (OK, but not very accurate)
Back azimuth $=252^{\circ} 10^{\prime} 53.2521^{\prime \prime}$
Convergence $=\left(252^{\circ} 10^{\prime} 53.2521^{\prime \prime}\right)-\left(72^{\circ} 10^{\prime} 50.3098^{\prime \prime}\right)-180^{\circ}=+2.9423^{\prime \prime} \quad$ (the same!)

## Exercise 2.3: An approximate method for computing ellipsoidal distance

This gives a method for computing an approximate ellipsoidal distance between two points with geodetic coordinates (latitude, longitude, and ellipsoid height). For the GRS-80, WGS-84, Clarke 1866, and most other Earth ellipsoids, note the following:

## Rules of Thumb

## 3 <br> $\mathbf{1}$ arc-second of latitude $\approx \mathbf{1 0 1} \mathbf{~ f t ~ ( a c c u r a t e ~ t o ~ w i t h i n ~ a b o u t ~} \pm 0.3 \mathrm{ft}$ in US) <br> $\mathbf{1} \mathbf{a r c}-$ second of longitude $\approx \mathbf{1 0 1} \mathbf{f t} \times \cos (l a t i t u d e) \quad$ (short by about 0.5 ft in US)

Based on these relationships, we can compute an approximate distance, to wit:
Equation 2.6 Approximate ellipsoidal distance between a pair of geodetic coordinates

$$
s \approx 101 \sqrt{\left(\Delta \varphi^{\prime \prime}\right)^{2}+(\Delta \lambda " \cos \bar{\varphi})^{2}} \text { feet }
$$

This equation is accurate to within about $\pm 1 \%$ everywhere on the Earth (and about $\pm 0.5 \%$ in AZ) where $\Delta \varphi^{\prime \prime}$ is change in latitude between two points in arc-seconds
$\Delta \lambda^{\prime \prime}$ is change in longitude between two points in arc-seconds
$\bar{\varphi}$ is average latitude of the two points

## Example computation

Given: Points CAS-2 and CAS-3 from the previous example (in Exercise 2.2).
Find: The approximate ellipsoidal distance between the points CAS-2 and CAS-3.

## Computations:

From the previous example, the average latitude of CAS-2 and CAS-3 is $\bar{\varphi}=34.5497141306^{\circ}$

$$
s \approx 101 \sqrt{\left(\Delta \varphi^{\prime \prime}\right)^{2}+\left(\Delta \lambda^{\prime \prime} \cos \bar{\varphi}\right)^{2}}
$$

$$
=101 \times \sqrt{\left.\left(\_^{-}-\quad\right)^{2}+[(\ldots)) \times \cos \left(34.5498030194^{\circ}\right)\right]^{2}}
$$

$$
=101 \times \sqrt{(\square)^{2}+(\square)^{2}}=
$$

$\qquad$
Solution:

$$
\begin{aligned}
&\left.\left.=101 \times \sqrt{(59.98077-58.60097)^{2}+[(42.59198}-47.78016\right) \times \cos \left(34.5498030194^{\circ}\right)\right]^{2} \\
&= 101 \times \sqrt{(1.37980)^{2}+(-4.27316)^{2}}=\underline{\mathbf{4 5 4} \mathbf{f t}}
\end{aligned}
$$

## Check using NGS Inverse tool:

Actual ellipsoid distance $($ geodesic $)=138.9428 \mathrm{~m}=\underline{455.849 \mathrm{ift}}$
Approximate geodetic inverse error $=-1.8 \mathrm{ft}=-0.4 \%$

## Exercise 2.4: A more accurate method for approximating ellipsoidal distance

Computation of accurate geodetic distances is difficult, but a good approximation over short distances can be computed using spherical angles based on an appropriate radius of curvature.
Equation 2.7 Central angle between two points on surface of a sphere

$$
\psi=\cos ^{-1}\left(\sin \varphi_{A} \sin \varphi_{B}+\cos \varphi_{A} \cos \varphi_{B} \cos \left(\lambda_{A}-\lambda_{B}\right)\right)
$$

where $\varphi_{A}, \varphi_{B}$ are the latitudes at points $A$ and $B$, respectively
$\lambda_{A}, \lambda_{B}$ are longitudes at points $A$ and $B$, respectively

Equation 2.8 Approximate geodetic inverse based on spherical angle

$$
s=R_{\alpha} \psi=\left(\frac{R_{M} R_{N}}{R_{M} \sin ^{2} \widetilde{\alpha}_{A B}+R_{N} \cos ^{2} \tilde{\alpha}_{A B}}\right) \psi
$$

where all variables are as defined previously and radii of curvature are evaluated at the mean latitude of the two points.

The accuracy of the distances computed by Equation 2.8 vary with azimuth, and are generally shorter than actual by a maximum of 10 ppm for distances less of than about 10 miles (e.g., a one mile inverse is at most 0.05 ft shorter than actual).
A highly accurate method for computing geodetic distance and azimuth was published by Vincenty (1975), and is the one used in the NGS geodetic tool "Inverse".

## Example computation

Given: Points CAS-2 and CAS-3 from the previous two examples (in Exercises 2.2 and 2.3).
Find: The approximate ellipsoidal distance between the points CAS-2 and CAS-3.

## Computations:

First compute the spherical angle,

$$
\begin{aligned}
& \psi=\cos ^{-1}\left(\sin \varphi_{A} \sin \varphi_{B}+\cos \varphi_{A} \cos \varphi_{B} \cos \left(\lambda_{A}-\lambda_{B}\right)\right) \\
& \psi=\cos ^{-1}[\sin (\square) \times \sin ( \\
& +\cos (\square) \times \cos (\square) \\
& \times \cos (\ldots)-(\ldots)] \\
& =\cos ^{-1}[(\square) \\
& +( \\
& \times \\
& \times \\
& =\cos ^{-1}[\square]=\quad{ }^{\circ}=\quad=
\end{aligned}
$$

Then compute the ellipsoid distance as (with $\psi$ in radians) as

$$
s=R_{\alpha} \psi=\left(\frac{R_{M} R_{N}}{R_{M} \sin ^{2} \widetilde{\alpha}_{A B}+R_{N} \cos ^{2} \widetilde{\alpha}_{A B}}\right) \psi .
$$

From Exercise 1.3, the radius of curvature is $R_{\alpha}=20,939,171 \mathrm{ift}$ (rounded to the nearest foot) at a mean latitude of $\bar{\varphi}=34.5498030194^{\circ}$ for points CAS-2 and CAS-3. This is at an (approximate) azimuth of $\widetilde{\alpha}=72.10473^{\circ}$ from CAS-2 to CAS-3 (Exercises 1.3 and 2.2).

The spherical angle must be converted to radians for this computation, as follows:

$$
\begin{array}{ccc}
s= & \times & R_{\alpha} \\
s= & \times & \circ \\
& & \\
& \\
180^{\circ} \\
& \\
\hline
\end{array}
$$

Solution:

$$
\begin{aligned}
& \psi=\cos ^{-1}\left(\sin \varphi_{A} \sin \varphi_{B}+\cos \varphi_{A} \cos \varphi_{B} \cos \left(\lambda_{A}-\lambda_{B}\right)\right) \\
& \psi=\cos ^{-1}\left[\sin \left(34.5496113805556^{\circ}\right) \times \sin \left(34.5499946583333^{\circ}\right)\right. \\
& +\cos \left(34.5496113805556^{\circ}\right) \times \cos \left(34.5499946583333^{\circ}\right) \\
& \left.\times \cos \left(-112.446605600000^{\circ}+112.445164438889^{\circ}\right)\right] \\
& =\cos ^{-1}[(0.567119620484369 \times 0.567125130147320) \\
& +(0.823635438808740 \times 0.823631645066765 \times 0.999999999683663)] \\
& =\cos ^{-1}[0.321627788576296+0.678372211186735]=\underline{0.0012473344^{\circ}}=\underline{4.490404^{\prime}}
\end{aligned}
$$

The spherical angle must be converted to radians for this computation, as follows:

$$
s=20,939,171 \mathrm{ift} \times 0.0012473344^{\circ} \times \frac{3.14159265}{180^{\circ}}=\underline{\mathbf{4 5 5 . 8 4 8} \mathbf{~ i f t}}
$$

From Exercise 2.3, Vincenty inverse is nearly identical, $s=\underline{455.849 \mathrm{ift}}$

$$
\text { (error }=-0.001 \mathrm{ft}=-0.0002 \%)
$$

The results shown here were computed using Microsoft Excel, which has a numerical precision of 15 digits. Note that most hand calculators have difficulty accurately performing these calculations due to lower numerical precision. Example computations using different numerical precisions are given below (these will vary depending on the calculator, sequence of computations, and number of digits entered):

14 digits of numerical precision $\rightarrow s=455.845$ ift ( $-0.0009 \%$ error $)$
13 digits of numerical precision $\rightarrow s=455.821 \mathrm{ift}(-0.0061 \%$ error)
12 digits of numerical precision $\rightarrow s=455.611 \mathrm{ift}(-0.0522 \%$ error $)$

## Exercise 2.5: Deflection of the vertical and the Laplace correction

In general, the plumbline (gravity vector) passing through the axis of an instrument is not parallel to a line perpendicular to the reference ellipsoid (the ellipsoid normal), and the angle between these two lines is called the deflection of the vertical. The deflection of the vertical is divided into north-south and east-west components, denoted as $\xi$ and $\eta$, respectively. These can be obtained from the NGS model DEFLEC09 and USDOV2009 for any location in the US. DEFLEC09 was derived from the GEOID09 "hybrid" geoid model, and is the appropriate one to use for survey observations referenced to NAD 83. USDOV2009 was derived from the purely gravimetric geoid model USGG2009, which is referenced to ITRF 00.
If the deflection of the vertical is not zero, an instrument leveled to the local plumbline will not be "level" with respect to the geodetic datum. When using terrestrial (optical) instruments, this affects determination of coordinates and azimuths using astronomic (or gyroscopic) methods; reductions of terrestrial observations to the ellipsoid; and change in ellipsoid height. In addition, since deflection of the vertical varies with location, it can cause horizontal and vertical errors in terrestrial surveys that are similar to the misclosure that occurs if a traverse is performed with an improperly leveled instrument.
The Laplace correction is the difference between astronomic and geodetic azimuth caused by deflection of the vertical. A simplified version of the Laplace correction is given on NGS datasheets, and adding this value to (clockwise) astronomic azimuths will give the geodetic azimuth for an approximately horizontal line of sight between stations.

Equation 2.9 The simplified (horizontal) Laplace correction (assumes approximately horizontal line of sight, a clockwise positive azimuth, and a positive east deflection of the vertical):
$L=\alpha-A=-\eta \tan \varphi$
where $\alpha$ and $A$ are the geodetic and astronomic azimuths, respectively
$\eta$ is the deflection of the vertical component in the east-west (prime vertical) direction $\varphi$ is the geodetic latitude

## Rules of Thumb

Maximum deflection of the vertical in Arizona $=35$ arc-seconds (DEFLEC09)
3
Maximum Laplace correction magnitude in Arizona $=25$ arc-seconds (DEFLEC09)
Simplified Laplace correction error is less than approximately $10 \%$ for zenith angles within about $5^{\circ}$ of horizontal

## Example computation

Given: In Elbow Canyon of the Virgin Mountains of northwestern Arizona, GPS was used to locate the southwest corner and the west quarter corner of Section 16, T 39 N, R 15 W , Gila and Salt River Baseline and Meridian. The following NAD 83 coordinates were obtained:

| Station | Latitude | Longitude | Ellipsoid height |
| :--- | :---: | :---: | :---: |
| SW Corner S16 | $36^{\circ} 46^{\prime} 31.612844^{\prime} \mathrm{N}$ | $113^{\circ} 55^{\prime} 21.70113^{\prime \prime} \mathrm{W}$ | 3530.589 ift |
| W 1/4 Corner S16 | $36^{\circ} 46^{\prime} 57.75891$ "N | $113^{\circ} 55^{\prime} 21.69613^{\prime \prime} \mathrm{W}$ | 3275.291 ift |

The geodetic azimuth and horizontal ground distance from the southwest corner to the west quarter corner based on these coordinates is $0^{\circ} 00^{\prime} 31.73^{\prime \prime}$ and 2644.715 ift .

Find: The astronomic quadrant bearing from the southwest corner to the west quarter corner of Section 16.

## Computations:

For the southwest corner of Section 16, DEFLEC09 gives $\xi=9.18$ ", $\eta=-26.48$ ", and $L=$ 19.79". Equation 2.9 can be rearranged to compute the astronomic azimuth:
$A=\alpha-L=$ $\qquad$ - $\qquad$ $=$ $\qquad$
Astronomic quadrant bearing $=$ $\qquad$ (rounded to nearest arc-second)

## Solution:

For the southwest corner of Section 16, DEFLEC09 gives $\xi=9.18^{\prime \prime}, \eta=-26.48^{\prime \prime}$, and $L=$ 19.79". Equation 2.9 can be rearranged to compute the astronomic azimuth:

$$
A=\alpha-L=-0^{\circ} 00^{\prime} 31.73^{\prime \prime}-19.18^{\prime \prime}=0^{\circ} 00^{\prime} 12.55^{\prime \prime}
$$

Astronomic quadrant bearing $=\mathbf{N} \mathbf{0 0}{ }^{\circ} \mathbf{0 0}{ }^{\prime} \mathbf{1 3 "} \mathbf{E}$ (rounded to nearest arc-second)
Check: $L=-\eta \tan \varphi=-\left(+26.48^{\prime \prime}\right) \times \tan \left(36^{\circ} 46^{\prime} 32^{\prime} \mathrm{N}\right)=19.79$ ", as given by DEFLEC09.

## How accurate is the simplified (horizontal) Laplace correction?

The complete Laplace correction is given by $L=-\eta \tan \varphi-(\xi \sin \alpha-\eta \cos \alpha) \cot \zeta$, where the first term is the same as Equation 2.9, and the second term is referred to as the deflection correction. The quantity $\zeta$ is the geodetic zenith angle, which can be estimated using the ellipsoid height difference and distance between the corners. Earth curvature increases the zenith angle, and can be accounted for by subtracting $0.0239 \times$ (distance in thousands of feet) ${ }^{2}$ from the height difference (this correction is covered in more detail in Exercise 4.2):

$$
\zeta=90^{\circ}-\tan ^{-1}\left\{\left[(3275.291-3530.589)-0.0239 \times 2.644715^{2}\right] / 2644.715\right\}=95.517^{\circ} .
$$

Thus the deflection correction is:

$$
\left[9.18^{\prime \prime} \times \sin \left(0^{\circ} 00^{\prime} 31.73^{\prime \prime}\right)-\left(-26.48^{\prime \prime}\right) \times \cos \left(0^{\circ} 00^{\prime} 31.73^{\prime \prime}\right)\right] \times \cot \left(95.517^{\circ}\right)=\underline{-2.56^{\prime \prime}}
$$

This gives a complete Laplace correction of $L=19.79$ " $-(-2.56 ")=22.35 "$. Although this deflection correction is rather large, note that this is a worse-case scenario, because the deflection of the vertical value in this example is essentially the maximum for Arizona. In most cases, the deflection correction is smaller than can be resolved using optical methods, and the simplified Laplace correction will suffice. This helps tremendously, since the simplified Laplace correction does not depend on the azimuth or zenith angle between stations, and so a unique value can be specified at a point.

## Section 3

## GRID COORDINATE SYSTEMS AND COMPUTATIONS

## How are the data displayed? How are the data used?

## Examples of grid coordinate errors for Arizona

Table 3.1 Examples of various positioning error sources and their magnitudes for Arizona due to grid coordinate system and computation problems (abbreviations and technical terms are defined in the Glossary).

| Positioning error examples for Arizona | Error magnitudes |
| :--- | :---: |
| Using SPCS 27 projection parameters for SPCS 83 projects | 37.9 miles (horizontal) |
| Determining State Plane coordinates in US Survey Feet when <br> International Feet are required | Up to 5 feet (horizontal) |
| Determining UTM coordinates in US Survey Feet when <br> International Feet are required | Up to 27 feet (horizontal) |
| Using linear coordinates from a geographic "projection" to <br> compute distances | Up to $\sim 1000$ feet horizontal per <br> mile |
| Using SPCS grid distances when "ground" distances are <br> required (example here is for Flagstaff) | $\sim 2.3$ feet horizontal per mile at <br> elevation of 7000 feet |
| Using UTM grid distances when "ground" distances are <br> required (example here is for Flagstaff) | $\sim 3.6$ feet horizontal per mile at <br> elevation of 7000 feet |
| Using planar computation methods to transform geodetically- <br> derived horizontal coordinates (example here is for | Varies, but increases rapidly <br> with size of area <br> converting from UTM to SPCS over a 20 mi $\times 20$ mi area in <br> (3 to 4 feet of horizontal <br> error for this example) |
| based on two common points) | area using planar scaling, rotation, and translation |

## Grid coordinate system information in NGS Datasheets and OPUS output

Both NGS Datasheets and OPUS output use the geodetic coordinates of the point to compute grid (map projection) coordinates in the State Plane and Universal Transverse Mercator coordinate systems. They also provide the convergence angle, grid point scale factor, and combined scale factor for both systems.
Portion of NGS Datasheet for station PR 23


Portion of OPUS output for station CAS-1


## Map projection distortion

Map projection distortion is an unavoidable consequence of attempting to represent a curved surface on a flat surface. It can be thought of as a change in the "true" relationship between points located on the surface of the Earth and the representation of their relationship on a plane. Distortion cannot be eliminated - it is a Fact of $\boldsymbol{L}$ ife. The best we can do is decrease the effect.

There are two general types of map projection distortion:

1. Linear distortion. Difference in distance between a pair of grid (map) coordinates when compared to the true ("ground") distance, denoted here by $\boldsymbol{\delta}$.

- Can express as a ratio of distortion length to ground length:
- E.g., feet of distortion per mile; parts per million (= mm per km)
- Note: 1 foot $/ \mathrm{mile}=189 \mathrm{ppm}=189 \mathrm{~mm} / \mathrm{km}$
- Linear distortion can be positive or negative:
- NEGATIVE distortion means the grid (map) length is SHORTER than the "true" horizontal (ground) length.
- POSITIVE distortion means the grid (map) length is LONGER than the "true" horizontal (ground) length.

2. Angular distortion. For conformal projections (e.g., Transverse Mercator, Lambert Conformal Conic, Stereographic, Oblique Mercator, etc.), it equals the convergence (mapping) angle, $\boldsymbol{\gamma}$. The convergence angle is the difference between grid (map) north and true (geodetic) north.

- Convergence angle is zero on the projection central meridian, positive east of the central meridian, and negative west of the central meridian
- Magnitude of the convergence angle increases with distance from the central meridian, and its rate of change increases with increasing latitude:

| Latitude | Convergence angle <br> 1 mile from CM | Latitude | Convergence angle <br> 1 mile from CM |
| :---: | :---: | :---: | :---: |
| $0^{\circ}$ | $0^{\circ} 00^{\prime} 00^{\prime \prime}$ | $50^{\circ}$ | $\pm 0^{\circ} 01^{\prime} 02^{\prime \prime}$ |
| $10^{\circ}$ | $\pm 0^{\circ} 00^{\prime} 09^{\prime \prime}$ | $60^{\circ}$ | $\pm 0^{\circ} 01^{\prime} 30^{\prime \prime}$ |
| $20^{\circ}$ | $\pm 0^{\circ} 00^{\prime} 19^{\prime \prime}$ | $70^{\circ}$ | $\pm 0^{\circ} 02^{\prime} 23^{\prime \prime}$ |
| $30^{\circ}$ | $\pm 0^{\circ} 00^{\prime} 30^{\prime \prime}$ | $80^{\circ}$ | $\pm 0^{\circ} 04^{\prime} 54^{\prime \prime}$ |
| $40^{\circ}$ | $\pm 0^{\circ} 00^{\prime} 44^{\prime \prime}$ | $89^{\circ}$ | $\pm 0^{\circ} 49^{\prime} 32^{\prime \prime}$ |

- Usually convergence is not as much of a concern as linear distortion, and it can only be minimized by staying close to the projection central meridian (or the Equator).

Total linear distortion of grid (map) coordinates is a combination of distortion due to Earth curvature and distortion due to ground height above the ellipsoid. In many areas, distortion due to variation in ground height is greater than that due to curvature. This is illustrated in the diagrams and tables on the following pages.

Table 3.2 Horizontal distortion of grid coordinates due to Earth curvature

| Maximum <br> zone width for <br> secant projections <br> (miles) | Maximum linear horizontal distortion, $\boldsymbol{\delta}$ |  |  |
| :---: | :---: | :---: | :---: |
|  | Parts per <br> million | Feet per mile | Ratio <br> (absolute value) |
| 16 miles | $\pm 1 \mathrm{ppm}$ | $\pm 0.005 \mathrm{ft} / \mathrm{mile}$ | $1: 1,000,000$ |
| 50 miles | $\pm 10 \mathrm{ppm}$ | $\pm 0.05 \mathrm{ft} / \mathrm{mile}$ | $1: 100,000$ |
| $\mathbf{7 1}$ miles | $\pm \mathbf{2 0} \mathbf{~ p p m}$ | $\pm \mathbf{0 . 1 ~ f t} / \mathrm{mile}$ | $\mathbf{1 : 5 0 , 0 0 0}$ |
| 112 miles | $\pm 50 \mathrm{ppm}$ | $\pm 0.3 \mathrm{ft} / \mathrm{mile}$ | $1: 20,000$ |
| 158 miles (e.g., SPCS) | $*$ | $\pm 100 \mathrm{ppm}$ | $\pm 0.5 \mathrm{ft} / \mathrm{mile}$ |
| 317 miles (e.g., UTM) ${ }^{\dagger}$ | $\pm 400 \mathrm{ppm}$ | $\pm 2.1 \mathrm{ft} / \mathrm{mile}$ | $1: 10,000$ |

*State Plane Coordinate System; zone width shown is valid between $\sim 0^{\circ}$ and $45^{\circ}$ latitude
${ }^{\dagger}$ Universal Transverse Mercator; zone width shown is valid between $\sim 30^{\circ}$ and $60^{\circ}$ latitude


Table 3.3 Horizontal distortion of grid coordinates due to ground height above the ellipsoid

| Height below (-) <br> and above (+) <br> projection surface | Maximum linear horizontal distortion, $\boldsymbol{\delta}$ |  |  |
| :---: | :---: | :---: | :---: |
|  | Parts per <br> million | Feet per mile | Ratio <br> (absolute value) |
| -100 feet, +100 feet | $\pm 4.8 \mathrm{ppm}$ | $\pm 0.03 \mathrm{ft} / \mathrm{mile}$ | $\sim 1: 209,000$ |
| $-\mathbf{4 0 0}$ feet, $+\mathbf{4 0 0}$ feet | $\pm \mathbf{1 9} \mathbf{~ p p m}$ | $\pm \mathbf{0 . 1} \mathrm{ft} / \mathrm{mile}$ | $\sim \mathbf{1}: \mathbf{5 2 , 0 0 0}$ |
| -1000 feet, +1000 feet | $\pm 48 \mathrm{ppm}$ | $\pm 0.3 \mathrm{ft} / \mathrm{mile}$ | $\sim 1: 21,000$ |
| +2000 feet | -96 ppm | $-0.5 \mathrm{ft} / \mathrm{mile}$ | $\sim 1: 10,500$ |
| +4000 feet* | -191 ppm | $-1.0 \mathrm{ft} / \mathrm{mile}$ | $\sim 1: 5200$ |
| +7000 feet** | -335 ppm | $-1.8 \mathrm{ft} / \mathrm{mile}$ | $\sim 1: 3000$ |
| $+12,600$ feet ${ }^{\dagger}$ | -603 ppm | $-3.2 \mathrm{ft} / \mathrm{mile}$ | $\sim 1: 1700$ |

*Approximate average topographic height in Arizona
** Approximate topographic height of Flagstaff, Arizona
${ }^{\dagger}$ Approximate maximum topographic height in Arizona

## $\beta^{\text {Rule of Thumb: }}$

A $100-\mathrm{ft}$ change in height causes a 4.8 ppm change in distortion


## Exercise 3.1: Distortion computations

Linear distortion is the ratio of grid distance to horizontal ground distance. One way to estimate distortion is to compute the distance between a pair of points based on the grid coordinates determined by the GPS software. This grid distance can then be divided by the ground distance between these points measured using a (properly calibrated) tape or EDM.

Equation 3.1 Approximating distortion at a point using measured grid and ground distances
$\delta \approx\left(\frac{\sqrt{\Delta N^{2}+\Delta E^{2}}}{\text { measured horizontal ground distance }}\right)-1$

Distortion can be computed more accurately (and conveniently) at a single point using the familiar "combined scale factor" approach:
Equation 3.2 Computing distortion at a point using Earth radius

$$
\delta=k\left(\frac{R_{G}}{R_{G}+h}\right)-1
$$

## Example computation

Given: Points CAS-2 and CAS-3 from the previous examples. The ellipsoid heights ( $h$ ) of these points are listed below, along with the grid coordinates and grid point scale factors $(k)$ derived from the adjusted geodetic coordinates (given in Exercise 2.2). A horizontal ground distance of 455.968 ift was carefully measured between these points.

CAS-2: NAD 83 (2002.0) ellipsoid height, $h=5466.883 \mathrm{ift}$

| Coordinate system | Northing, $\boldsymbol{N}$ (ift) | Easting, $\boldsymbol{E}$ (ift) | Grid scale factor, $\boldsymbol{k}$ |
| :---: | :---: | :---: | :---: |
| SPCS 83, Arizona Central (0202) | $1,291,805.295$ | $540,432.685$ | 0.999929147 |
| UTM 83, Zone 12 North | $12,546,092.208$ | $1,204,955.902$ | 0.999817145 |
| Low Distortion Projection (LDP) | $18,061.311$ | $56,042.621$ | 1.000258042 |

CAS-3: NAD 83 (2002.0) ellipsoid height, $h=5445.959 \mathrm{ift}$

| Coordinate system | Northing, $\boldsymbol{N}$ (ift) | Easting, $\boldsymbol{E}$ (ift) | Grid scale factor, $\boldsymbol{k}$ |
| :---: | :---: | :---: | :---: |
| SPCS 83, Arizona Central (0202) | $1,291,942.505$ | $540,867.361$ | 0.999928988 |
| UTM 83, Zone 12 North | $12,546,225.452$ | $1,205,391.755$ | 0.999816711 |
| Low Distortion Projection (LDP) | $18,200.930$ | $56,476.686$ | 1.000258048 |

Find: The linear distortion (in parts per million) at the midpoint between points CAS-2 and CAS-3 in SPCS, UTM, and LDP coordinates using both Equations 3.1 and 3.2 (the geometric mean radius of curvature $R_{G}=20,900,487 \mathrm{ift}$ was determined at the midpoint in Exercise 1.3).

Computations: For midpoint, use the mean grid scale factor and mean ellipsoid height $=5456 \mathrm{ft}$.

## SPCS 83 AZ C

Using Equation 3.1:

$\qquad$ $-1 \rightarrow$ in parts per million $\qquad$ $\times 1,000,000=$ $\qquad$

Using Equation 3.2:


## UTM 83 12N

Using Equation 3.1:


$$
=\ldots-1 \rightarrow \text { in parts per million } \rightarrow \ldots \times 1,000,000=
$$

Using Equation 3.2:

$\qquad$ $-1 \rightarrow$

## LDP

Using Equation 3.1:
$-1 \rightarrow$ in parts per million $\rightarrow$
$\times 1,000,000=$

Using Equation 3.2:
$\delta=\xlongequal[2]{+}\left(\frac{\square}{\square+\ldots}\right)-1=$ $\qquad$ $-1 \rightarrow$

Solution: For midpoint, use the mean grid scale factor and mean ellipsoid height $=5456 \mathrm{ft}$.

## SPCS 83 AZ C

## Using Equation 3.1:

$$
\begin{gathered}
\delta \approx\left(\frac{\sqrt{(1,291,942.505-1,291,805.295)^{2}+(540,867.361-540,432.685)^{2}}}{455.968}\right)-1=\left(\frac{455.817}{455.968}\right)-1 \\
=0.9996688-1 \rightarrow \text { in parts per million } \rightarrow-0.0003312 \times 1,000,000=\underline{\mathbf{- 3 3 1 . 2} \mathbf{~ p p m}}
\end{gathered}
$$

## Using Equation 3.2:

$$
\begin{array}{r}
\left.\delta=\frac{0.9999291+0.9999290}{2}\left(\frac{20,900,487}{20,900,487+5456}\right)-1=0.9996681-1 \rightarrow \frac{-\mathbf{3 3 1 . 9} \mathbf{~ p p m}}{(=-1.75 \mathrm{ft} / \mathrm{mile})}\right)
\end{array}
$$

## UTM 83 12N

Using Equation 3.1:

$$
\begin{gathered}
\delta \approx\left(\frac{\sqrt{(12,546,225.452-12,546,092.208)^{2}+(1,205,391.755-1,204,955.902)^{2}}}{455.968}\right)-1=\left(\frac{455.766}{455.968}\right)-1 \\
=0.9995570-1 \rightarrow \text { in parts per million } \rightarrow-0.0004430 \times 1,000,000=\underline{\mathbf{4 4 3 . 0} \mathbf{~ p p m}}
\end{gathered}
$$

Using Equation 3.2:

$$
\delta=\frac{0.9998171+0.9998167}{2}\left(\frac{20,900,487}{20,900,487+5456}\right)-1=0.9995560-1 \rightarrow \underline{-444.0} \mathbf{~ p p m}
$$

$$
(=-2.34 \mathrm{ft} / \mathrm{mile})
$$

## LDP

Using Equation 3.1:

$$
\delta \approx\left(\frac{\sqrt{(18,200.930-18,061.311)^{2}+(56,476.686-56,042.621)^{2}}}{455.968}\right)-1=\left(\frac{455.967}{455.968}\right)-1
$$

$$
=0.9999978-1 \rightarrow \text { in parts per million } \rightarrow-0.0000022 \times 1,000,000=\underline{-\mathbf{2} .2} \mathbf{~ p p m}
$$

Using Equation 3.2:

$$
\begin{equation*}
\delta=\frac{1.00025804+1.00025805}{2}\left(\frac{20,900,487}{20,900,487+5456}\right)-1=0.9999970-1 \rightarrow \underline{\mathbf{- 3 . 0} \mathbf{p p m}} \tag{=-0.016ft/mile}
\end{equation*}
$$

## Exercise 3.2: Six steps for designing a low-distortion grid coordinate system

1. Define project area and choose representative ellipsoid height, $\boldsymbol{h}_{\mathbf{0}}$ (not elevation)

- The average height of an area may not be appropriate (e.g., project near a mountain)
- No need to estimate height to an accuracy of better than about $\pm 20$ feet
- Note that as the size of the area increases, the effect of Earth curvature on distortion increases and it must be considered in addition to the effect of topographic height
- E.g., for areas wider than about 35 miles (perpendicular to the projection axis), distortion due to curvature alone exceeds 5 parts per million (ppm)

2. Place central meridian near centroid of project area
3. Scale central meridian of projection to representative ground height, $\boldsymbol{h}_{\mathbf{0}}$

Equation 3.3 Local map projection scaled to "ground"

$$
k_{0}=1+\frac{h_{0}}{R_{G}}
$$

- Where $R_{G}$ is geometric mean radius of curvature, $R_{G}=\frac{a \sqrt{1-e^{2}}}{1-e^{2} \sin ^{2} \varphi} \quad$ (Equation 1.7)
- Alternatively, can initially approximate $R_{G}$ as $20,900,000$ feet for Arizona (since $k_{\mathrm{o}}$ will likely be refined in Step \#4)
- This procedure is for the Transverse Mercator projection
- For Lambert Conical projection, use same equation for scale of standard parallel

4. Check distortion at points distributed throughout project area

- Best approach is to compute distortion over entire area and generate distortion contours (this ensures optimal low-distortion coverage)
- May require repeated evaluation using different $k_{\mathrm{o}}$ values
- Distortion computed at a point as $\delta=k\left(\frac{R_{G}}{R_{G}+h}\right)-1 \quad$ (Equation 3.2)
- Where $k=$ projection grid scale factor at a point (with respect to ellipsoid; see Equations 3.3 and 3.4)
- Multiply $\delta$ by $1,000,000$ to get distortion in parts per million ( ppm )

5. Keep the definition SIMPLE and CLEAN!

- Define $k_{\mathrm{o}}$ to no more than SIX decimal places, e.g., 1.000206 (exact)
- Note: A change of one unit in the sixth decimal place equals distortion caused by a 21-foot change in height
- Defining central meridian and latitude of grid origin to nearest whole arc-minute usually adequate (e.g., Central meridian $=111^{\circ} 48^{\prime} 00^{\prime \prime} \mathrm{W}$ )
- Define grid origin using large whole values with as few digits as possible (e.g., False easting $=50,000 ; \quad$ Max coordinate $<100,000$ )


## 6. Explicitly define linear unit and geodetic datum

- E.g., Linear unit = International foot; Geodetic datum = NAD 83(2007)


## Example computation

Design a Low Distortion Projection (LDP) for Prescott

1. Define project area and choose representative ellipsoid height, $\boldsymbol{h}_{\mathbf{0}}$ (not elevation)

From topographic maps and benchmark information, a representative ellipsoid height is $\boldsymbol{h}_{\mathbf{0}}=\underline{\mathbf{5 4 0 0 f t}}$ (no need for greater accuracy than nearest $\pm 10$ feet)
2. Place central meridian near centroid of project area

Based on location and extent of Prescott, a good, clean value is $\boldsymbol{\lambda}_{\mathbf{0}}=\underline{\mathbf{1 1 2}^{\circ} \mathbf{2 8}} \mathbf{} \mathbf{0 0} \mathbf{0 0} \mathbf{W}$
3. Scale central meridian of projection to representative ground height, $\boldsymbol{h}_{\mathbf{0}}$

First compute Earth radius at mid-latitude of Prescott, $\varphi=34^{\circ} 32^{\prime} 00^{\prime \prime} \mathrm{N}$ (no need for greater accuracy than nearest arc-minute of latitude):

$$
R_{G}=\frac{a \sqrt{1-e^{2}}}{1-e^{2} \sin ^{2} \varphi}=\frac{20,925,646.325 \times \sqrt{1-0.006694380023}}{1-0.006694380023 \times\left[\sin \left(34.533333^{\circ}\right)\right]^{2}}=\underline{20,900,450 \mathrm{ift}}
$$

Thus the central meridian scale factor scaled to the representative ellipsoid height is

$$
k_{0}=1+\frac{h_{0}}{R_{G}}=1+\frac{5400}{20,900,450}=\underline{\mathbf{1 . 0 0 0 2 5 8}}
$$

Based on these results, the following Transverse Mercator projection is defined (will refine definition if necessary based on results of Step \#4):

| Latitude of grid origin, | $\varphi_{0}=34^{\circ} 30^{\prime} 00^{\prime \prime} \mathrm{N}$ |
| :--- | :--- |
| Longitude of central meridian, | $\lambda_{0}=112^{\circ} 28^{\prime} 00^{\prime \prime} \mathrm{W}$ |
| False northing, | $N_{0}=0.000 \mathrm{ift}$ |
| False easting, | $E_{0}=50,000.000 \mathrm{ift}$ |
| Central meridian scale factor, | $k_{0}=1.000258$ |

## 4. Check distortion at points distributed throughout project area

Distortion can be computed at various points throughout the project area. These can be survey control points or even artificial points taken from topographic maps.
To illustrate, we can use the results of the point distortion computation CAS-2 from the previous example (which is repeated here for convenience)

$$
\delta=k\left(\frac{R_{G}}{R_{G}+h}\right)-1=1.00025804 \times\left(\frac{20,900,487}{20,900,487+5467}\right)-1=0.999996468-1=\underline{\mathbf{- 3 . 5} \mathbf{p p m}}
$$

For CAS-4 (at the hospital, 230 ft lower than the resort) we have:

$$
\delta=k\left(\frac{R_{G}}{R_{G}+h}\right)-1=1.00025802 \times\left(\frac{20,900,495}{20,900,495+5235}\right)-1=1.000007546-1=\underline{+\mathbf{7 . 5} \mathbf{p p m}}
$$

This computation can be performed at discrete points throughout the project area, but best approach is to compute distortion over entire area (for example on a 3-arc-second grid) and generate distortion contours to ensure optimal low-distortion coverage.

The ability to achieve low distortion is limited by change in elevation (height) within the project area. A reasonable goal might be to limit distortion to $\pm 0.1 \mathrm{ft}$ per mile, which is about $\pm 20 \mathrm{ppm}$ and corresponds to a height change of about $\pm 400 \mathrm{ft}$.
5. Keep the definition SIMPLE and CLEAN!

All of the projection parameters were initially defined in Step \#3, but trial-and-error may be necessary to refine definition.

- Note $k_{\mathrm{o}}$ is defined to exactly SIX decimal places: $\boldsymbol{k}_{\mathbf{0}}=\mathbf{1 . 0 0 0 2 5 8}$ (exact)
- Both latitude of grid origin and central meridian are defined to nearest whole arc-minute:

$$
\varphi_{0}=34^{\circ} 30^{\prime} \mathbf{0 0} " \mathrm{~N} \text { and } \lambda_{0}=112^{\circ} 28^{\prime} \mathbf{0 0} " \mathrm{~W}
$$

$\varphi_{0}$ was selected far enough south to ensure positive northings, but far enough north to keep northings less than $100,000 \mathrm{ift}$.

- Grid origin is defined using clean whole values with as few digits as possible:

$$
N_{0}=0.000 \text { ift } \quad \text { and } \quad E_{0}=50,000.000 \text { ift }
$$

These values were selected to keep grid coordinates positive but less than 100,000 ift within the Prescott area (it is conventional to set $N_{0}$ to zero at $\varphi_{0}$, but is not required).

## 6. Explicitly define linear unit and geodetic datum

Linear unit is International foot, and geodetic datum is NAD 83(2007)
Final Low Distortion Projection definition for this example:
Linear unit: International foot
Geodetic datum: North American Datum of 1983(2007)
System: Arizona LDP
Zone: Prescott
Projection: Transverse Mercator
Latitude of grid origin: $34^{\circ} 30^{\prime} 00^{\prime \prime} \mathrm{N}$
Longitude of central meridian: $112^{\circ} 28^{\prime} 00^{\prime \prime} \mathrm{W}$
Northing at grid origin: 0.000 ft
Easting at central meridian: $50,000.000 \mathrm{ft}$
Scale factor on central meridian: 1.000258 (exact)
Note that this coordinate system definition only deals with horizontal coordinates (no vertical datum is specified).

## Methods for creating low-distortion grid coordinate systems

1. Design a Low Distortion Projection (LDP) for a specific project geographic area.

Use a conformal projection referenced to the existing geodetic datum.
Described in detail previously in this document.
2. Scale the reference ellipsoid "to ground".

A map projection referenced to this new "datum" is then designed for the project area.

## Problems:

- Requires a new ellipsoid (datum) for every coordinate system, which makes it more difficult to implement than an LDP.
- New datum makes it more complex than an LDP, yet it does not perform any better.
- Generates new set of latitudes that can be substantially different from original latitudes.
- Change in latitude can exceed 3 feet per 1000 ft of topographic height, depending on method used for scaling the ellipsoid (this case is for scaling with constant flattening).
- Can lead to confusion over which latitude values are correct.

3. Scale an existing published map projection "to ground".

Referred to as "modified" State Plane when an existing SPCS projection definition is used.

## Problems:

- Generates coordinates with values similar to "true" State Plane (can cause confusion).
- Can eliminate this problem by translating grid coordinates to get smaller values.
- Often yields "messy" parameters when a projection definition is back-calculated from the scaled coordinates (in order to import the data into a GIS).
- More difficult to implement in a GIS, and may cause problems due to rounding or truncating of "messy" projection parameters (especially for large coordinate values).
- Can reduce this problem through judicious selection of "scaling" parameters.
- Does not reduce the convergence angle (it is same as that of original SPCS definition).
- In addition, the arc-to-chord correction may be significant; it can reach $1 / 2 \mathrm{arc}$-second for a 1 -mile line located 75 miles from the projection axis (this correction is used along with the convergence angle for converting grid azimuths to geodetic azimuths).
- MOST IMPORTANT: Usually does not minimize distortion over as large an area as the other two methods.
- Extent of low-distortion coverage generally decreases as distance increases from projection axis (i.e., central meridian for TM and central parallel for LCC projection).
- State Plane axis usually does NOT pass through the project area.
- Sketches illustrating this problem with "modified" SPCS are shown on the next page.



## Exercise 3.3: Two methods for computing horizontal "ground" distance

This exercise gives two simple methods for computing horizontal "ground" distances using geodetic information. The first method is done by scaling the ellipsoid distance (geodesic) using the average of the ellipsoid heights at the endpoints, as follows:

Equation 3.4 Approximate geodetic "ground" distance based on ellipsoid distance (geodesic)

$$
D_{g r n d}=s\left(1+\frac{\bar{h}}{\bar{R}_{G}}\right)
$$

where $s$ is the ellipsoid distance (geodesic)
$\bar{h}$ is the average ellipsoid height of the two points
$\bar{R}_{G}$ is the geometric mean radius of curvature at the midpoint latitude of the two points
The second method for computing a horizontal ground distance can be done by using a GPS (GNSS) vector directly. Neglecting Earth curvature, this distance can be computed as:

Equation 3.5 Approximate "ground" distance based on GPS (GNSS) vector components

$$
D_{g r n d}=\sqrt{\Delta X^{2}+\Delta Y^{2}+\Delta Z^{2}-\Delta h^{2}}
$$

where $\Delta X, \Delta Y, \Delta Z$ are the GPS vector components (as ECEF Cartesian coordinate deltas)
$\Delta h=$ change in ellipsoid height between vector end points
Note that this method can also be used with end point coordinates (rather than a GPS vector), by converting the latitude, longitude, and ellipsoid heights to $X, Y, Z$ ECEF coordinates using Equation 2.1, and then using the difference in ECEF coordinates in Equation 3.5.

Accounting for curvature increases this horizontal ground distance, but for distances of less than 20 miles (about 30 km ), the increase is less than 1 part per million ( ppm ), i.e., less than 0.1 ft ( 3 cm ). The horizontal distance can be multiplied by the following curvature correction factor to get the approximate curved horizontal ground distance (error is less than about 0.01 ft for distances under 50 miles):

Equation 3.6 Correction factor applied to horizontal distance to account for curvature

$$
C_{C}=\frac{2 \bar{R}_{G} \sin ^{-1}\left(\frac{D_{\text {grrd }}}{2 \bar{R}_{G}}\right)}{D_{\text {grnd }}}
$$

where all variables are as defined previously. An Earth radius of $20,900,000 \mathrm{ft}$ is sufficiently accurate in Arizona for distances of less than about 100 miles (causes less than 0.01 ft error).

## Example computation

Given: Points CAS-2 and CAS-3 from the previous exercises, and a GPS vector from CAS-2 to CAS-3 with components $\Delta X=438.001 \mathrm{ft}, \Delta Y=-76.678 \mathrm{ft}$, and $\Delta Z=103.056 \mathrm{ft}$.

Find: The horizontal "ground" distance between these points using the two methods in this exercise.

## Computations:

Method 1. From Exercises 2.3 and 2.4, ellipsoid distance (geodesic) is $s=455.849 \mathrm{ift}$
From Exercises 1.3 and 3.1, $R_{G}=20,900,487 \mathrm{ift}$ at midpoint between CAS-2 and CAS-3 (which is the same as the average $R_{G}$ for the two points)

From the ellipsoid heights in Exercise 3.1, the average ellipsoid height is

$$
\bar{h}=\left(\varlimsup^{+}\right) / 2=\ldots \text { ift }
$$

So ground distance is

$$
D_{\text {grnd }}=s\left(1+\frac{\bar{h}}{\bar{R}_{G}}\right)=\square \times(1+\square)=\square \quad \text { ift }
$$

Method 2. Using the given GPS vector components and $\Delta h$ from Exercises 3.1 gives a horizontal ground distance of

$$
D_{g r n d}=\sqrt{(\square)^{2}+(\square)^{2}+(\square)^{2}-(\square)^{2}}=\underline{\text { ift }}
$$

Solution:
Method 1. From Exercise 2.4, ellipsoid distance (geodesic) is $s=455.849 \mathrm{ift}$
From Exercises 1.3 and 3.1, $R_{G}=20,900,487 \mathrm{ift}$ at midpoint between CAS-2 and CAS-3 (which is the same as the average $R_{G}$ for the two points)

From the ellipsoid heights in Exercise 3.1, the average ellipsoid height is

$$
\bar{h}=(5466.883+5445.959) / 2=5456.421 \mathrm{ift}
$$

So ground distance is

$$
D_{g r d}=455.849 \times\left(1+\frac{5456.421}{20,900,487}\right)=455.849 \times 1.00026107=\underline{455.968} \text { ift }
$$

Method 2. Using the given GPS vector components and $\Delta h=5445.959-5466.883=-20.924 \mathrm{ft}$ gives a horizontal ground distance of

$$
D_{g r n d}=\sqrt{(438.001)^{2}+(-76.678)^{2}+(103.056)^{2}-(-20.924)^{2}}=\underline{455.968} \mathbf{i f t},
$$

which is the same as that computed for Method 1. For such a short distance, curvature is completely negligible. This can be verified using Equation 3.6, which gives a curvature correction factor of $C_{C}=1.000000000020$, or 0.00002 ppm . As stated previously, the curvature correction factor is less than 1 ppm for distances of less than about 20 miles.

## Exercise 3.4: Projection grid scale factor and convergence angle computation

For the Transverse Mercator projection, the grid scale factor at a point can be computed as follows (modified from Stem, 1990, pp. 32-35):

Equation 3.7 Transverse Mercator projection grid scale factor formula
$k=k_{0}\left\{1+\frac{(\Delta \lambda \cos \varphi)^{2}}{2}\left(1+\frac{e^{2} \cos ^{2} \varphi}{1-e^{2}}\right)\left[1+\frac{(\Delta \lambda \cos \varphi)^{2}}{12}\left(5-4 \tan ^{2} \varphi+\frac{e^{2} \cos ^{2} \varphi}{1-e^{2}}\left(9-24 \tan ^{2} \varphi\right)\right)\right]\right\}$
where $\Delta \lambda=\lambda_{0}-\lambda$ (in radians, for negative west longitude)
$\lambda=$ geodetic longitude of point
$\lambda_{0}=$ central meridian longitude
and all other variables are as defined previously.
The following shorter equation can be used to approximate $k$ for the Transverse Mercator projection. It is accurate to better than 0.02 part per million (at least 7 decimal places) if the computation point is within about $\pm 1^{\circ}$ of the central meridian (about 50 to 60 miles between latitudes of $30^{\circ}$ and $45^{\circ}$ ):

Equation 3.8 Approximate Transverse Mercator projection grid scale factor formula

$$
k \approx k_{0}\left\{1+\frac{(\Delta \lambda \cos \varphi)^{2}}{2}\left(1+\frac{e^{2} \cos ^{2} \varphi}{1-e^{2}}\right)\right\}
$$

Note that this equation may not be sufficiently accurate for computing $k$ throughout a UTM system zone (at the zone width of $\pm 3^{\circ}$ from the central meridian the error can exceed 1 ppm ).

An even simpler equation can be used to approximate the grid scale factor, which utilizes the grid coordinate easting value and is about twice as accurate as the previous equation (i.e., better than 0.01 part per million if the computation point is within about $\pm 1^{\circ}$ of the central meridian):

Equation 3.9 Another approximate Transverse Mercator projection grid scale factor formula

$$
k \approx k_{0}+\frac{\left(E_{0}-E\right)^{2}}{2\left(k_{0} R_{G}\right)^{2}}
$$

where $E=$ Easting of the point where $k$ is computed (in same units as $R_{G}$ )
$E_{0}=$ False easting (on central meridian) of projection definition (in same units as $R_{G}$ )
$R_{G}=$ Earth geometric mean radius of curvature (can use 20,900,000 feet for Arizona)

For the Lambert Conformal Conic projection, the grid scale factor at a point can be computed as follows (modified from Stem, 1990, pp. 26-29):

Equation 3.10 Lambert Conformal Conic projection grid scale factor formula $k=k_{0} \frac{\cos \varphi_{0}}{\cos \varphi} \sqrt{\frac{1-e^{2} \sin ^{2} \varphi}{1-e^{2} \sin ^{2} \varphi_{0}}} \exp \left\{\frac{\left.\left.\sin \varphi_{0}\left[\ln \frac{1+\sin \varphi_{0}}{1-\sin \varphi_{0}}-\ln \frac{1+\sin \varphi}{1-\sin \varphi}+e\left(\ln \frac{1+e \sin \varphi}{1-e \sin \varphi}-\ln \frac{1+e \sin \varphi_{0}}{1-e \sin \varphi_{0}}\right)\right]\right\}\right\}}{\}}\right.$
where $k_{0}=$ projection grid scale factor applied to central parallel (tangent to ellipsoid if $k_{0}=1$ ) $\varphi_{0}=$ geodetic latitude of central parallel = standard parallel for one-parallel LCC $e=\sqrt{e^{2}}=\sqrt{2 f-f^{2}}=$ first eccentricity of the reference ellipsoid
and all other variables are as defined previously. In order to use this equation for a two-parallel LCC, the two-parallel LCC must first be converted to an equivalent one-parallel LCC by computing $\varphi_{0}$ and $k_{0}$. The equations to do this are long, but are provided here for the sake of completeness. For a two-parallel LCC, the central parallel is

$$
\varphi_{0}=\sin ^{-1}\left[\frac{2 \ln \left(\frac{\cos \varphi_{\mathrm{S}}}{\cos \varphi_{\mathrm{N}}} \sqrt{\frac{1-e^{2} \sin ^{2} \varphi_{\mathrm{N}}}{1-e^{2} \sin ^{2} \varphi_{\mathrm{S}}}}\right)}{\ln \left(\frac{1+\sin \varphi_{\mathrm{N}}}{1-\sin \varphi_{\mathrm{N}}}\right)-\ln \left(\frac{1+\sin \varphi_{\mathrm{S}}}{1-\sin \varphi_{\mathrm{S}}}\right)+e\left[\ln \left(\frac{1+e \sin \varphi_{\mathrm{S}}}{1-e \sin \varphi_{\mathrm{S}}}\right)-\ln \left(\frac{1+e \sin \varphi_{\mathrm{N}}}{1-e \sin \varphi_{\mathrm{N}}}\right)\right]}\right],
$$

and the central parallel scale factor is

$$
k_{0}=\frac{\cos \varphi_{\mathrm{N}}}{\cos \varphi_{0}} \sqrt{\frac{1-e^{2} \sin ^{2} \varphi_{0}}{1-e^{2} \sin ^{2} \varphi_{\mathrm{N}}}}
$$

$$
\times \exp \left\{\frac{\sin \varphi_{0}}{2}\left[\ln \left(\frac{1+\sin \varphi_{\mathrm{N}}}{1-\sin \varphi_{\mathrm{N}}}\right)-\ln \left(\frac{1+\sin \varphi_{0}}{1-\sin \varphi_{0}}\right)+e\left(\ln \left[\frac{1+e \sin \varphi_{0}}{1-e \sin \varphi_{0}}\right]-\ln \left[\frac{1+e \sin \varphi_{\mathrm{N}}}{1-e \sin \varphi_{\mathrm{N}}}\right]\right)\right]\right\},
$$

where $\varphi_{\mathrm{N}}$ and $\varphi_{\mathrm{S}}$ = geodetic latitude of northern and southern standard parallels, respectively, and all other variables are as defined previously.

Convergence angles. For the TM , the convergence angle can be approximated as $\gamma=-\Delta \lambda \sin \varphi$ (where all variables are as defined previously; the units of $\gamma$ are the same as the units of $\Delta \lambda$ ). This equation is accurate to better than $\pm 00.2$ " if the computation point is within $\sim 1^{\circ}$ of the central meridian. For any LCC, the convergence angle is exactly $\gamma=-\Delta \lambda \sin \varphi_{0}$.

## Exercise 3.5: Grid versus geodetic bearings

Illustrates misclosure problem with geodetic azimuths, and shows how to convert grid azimuths to geodetic azimuths.

Equation 3.11 Relationship between grid and forward geodetic azimuth from point $A$ to $B$

$$
\alpha_{A B}=t_{A B}+\gamma_{A}-(t-T)_{A B}
$$

where $\alpha_{A B}$ and $t_{A B}=$ geodetic and grid azimuths from point $A$ to $B$, respectively
$\gamma_{A}=$ map projection convergence angle at point $A$
$(t-T)_{A B}=$ Arc-to-chord ("second term") correction from $A$ to $B$ (usually negligible)

## Example using Low Distortion Projection (LDP), State Plane, and UTM coordinates

Consider closed polygon below from points CAS-2 to CAS-3 to CAS-4 to CAS-2 (not to scale). Label the figure with distances, grid azimuths, and geodetic forward and back azimuths.


Example solution: Computed using Low Distortion Projection (LDP) coordinates

| MISCLOSURES (computed using LDP coordinates) |  |
| :--- | :--- |
| Grid bearings and grid distances (misclosure due to rounding) | 0.0007 ft |
| Grid bearings and "ground" distances | 0.0152 ft |
| Forward geodetic bearings and grid distances | 0.1617 ft |
| Forward geodetic bearings and "ground" distances | 0.1638 ft |
| Back geodetic bearings and grid distances | 0.1484 ft |
| Back geodetic bearings and "ground" distances | 0.1499 ft |
| Mean forward \& back geodetic bearings and grid distances | 0.0136 ft |
| Mean forward \& back geodetic bearings and "ground" distances | 0.0273 ft |

CAS-4

$$
\begin{gathered}
\gamma=+00^{\prime} 29.0^{\prime \prime}(\text { LDP }) \\
\gamma=-18^{\prime} 14.1 "(\text { SPCS }) \\
\gamma=-49^{\prime} 26.1 "(\text { UTM }
\end{gathered}
$$

## Notes

1) Misclosures the same for all grid coordinates systems.
2) Maximum magnitude of arc-to-chord correction $(t-T)$ :
a) $0.0015^{\prime \prime}$ for LDP coordinates
b) $0.0480^{\prime \prime}$ for SPCS 83 AZ C coordinates
c) 0.1321 " for UTM 8312 N coordinates

CAS-2

$$
\gamma=+00^{\prime} 41.0^{\prime \prime}(\text { LDP })
$$

$$
\gamma=-18 \prime 02.0^{\prime \prime}(\text { SPCS })
$$

$$
\gamma=-49^{\prime} 13.9^{\prime \prime}(\text { UTM })
$$

## Section 4

## VERTICAL DATUMS AND HEIGHT SYSTEMS

## How high is it? How deep is it? Where will water go?

## Examples of height determination errors for Arizona

Table 4.1 Examples of various positioning error sources and their magnitudes for Arizona due to vertical datum and height system problems (abbreviations and technical terms are defined in the Glossary).

| Positioning error examples for Arizona | Error magnitudes |
| :--- | :---: |
| Using NGVD 29 when NAVD 88 required | 1.2 to 4.5 feet (vertical) |
| Using ellipsoid heights for elevations | Varies from 63 feet to <br> 113 feet (vertical) |
| Neglecting geoid slope when transferring elevations with <br> GPS | Up to 0.7 foot vertical <br> per mile horizontal |
| Using geoid model GEOID03 when GEOID09 is required to <br> derive elevations from ellipsoid heights | Varies from -0.7 foot to <br> +0.5 foot (vertical) |
| Using leveling without orthometric corrections to "correct" <br> GPS-derived elevations | Can exceed 0.05 foot vertical <br> per mile horizontal |
| Generating GPS-derived elevations using a best-fit inclined <br> planar correction surface based on ties to inappropriate or <br> inconsistent vertical control (via a vertical "calibration" or <br> "localization") | Varies, but can cause very large <br> systematic vertical errors <br> (can exceed several feet) |

## Exercise 4.1: Ellipsoid, orthometric, and geoid heights

The relationship between ellipsoidal, orthometric, and geoid heights is shown in the figure below. Note that everywhere in the coterminous US, the geoid height is negative (i.e., the geoid is below the ellipsoid). But in most of Alaska, the geoid height is positive.


Equation 4.1 Relationship between ellipsoidal, orthometric, and geoid heights

$$
h=H+N_{G}
$$

where $h, H$, and $N_{G}$ are the ellipsoidal, orthometric, and geoid heights, respectively.
Strictly speaking, the relationship in Equation 4.1 is approximate due to deflection of the vertical. However, it is accurate at the sub-millimeter level, and so can be considered exact for all practical purposes.

## Rules of Thumb:

Accuracy of NAVD 88 orthometric heights derived from NAD 83 ellipsoid heights using the following geoid models (based on NGS documentation and given at the $95 \%$ confidence level):

GEOID09 approximate absolute accuracy: $\pm 0.09 \mathrm{ft}( \pm 2.7 \mathrm{~cm})$
GEOID03 approximate absolute accuracy: $\pm 0.15 \mathrm{ft}( \pm 4.7 \mathrm{~cm})$
GEOID99 approximate absolute accuracy: $\pm 0.30 \mathrm{ft}( \pm 9.0 \mathrm{~cm})$
GEOID96 approximate absolute accuracy: $\pm 0.35 \mathrm{ft}( \pm 10.8 \mathrm{~cm})$
The relative accuracy of these geoid models is 1 to 2 ppm , or better.

## Example computation

Given: An NGS Datasheet for conventional NGS control station PEND (below):


Find: The ellipsoid height of PEND in International and US Survey Feet.

## Computations:

Sometimes the only horizontal control station available for a GPS survey was determined using conventional methods. These do not have an ellipsoid height, but there is enough information to compute it if an accurate NAVD 88 orthometric height is available. From the Datasheet we have:

$$
\begin{array}{llll}
h= & H & + & N_{G} \\
h= & & +\ldots
\end{array}
$$

## Solution:


$h=2160.187 \mathrm{~m}+(-23.11 \mathrm{~m})=\underline{2137.08 \mathrm{~m}}=\underline{7011.42 \mathrm{ift}}=\underline{7011.40 \mathrm{sft}}$
 accuracy of the computation

## Exercise 4.2: Trigonometric leveling

Equation 4.2 Change in elevation from trigonometric leveling

$$
\Delta H=D_{S} \cos v+i-r+C_{C R}
$$

where $\Delta H$ is the change in elevation (nominally orthometric height)
$D_{S}$ is the slope distance
$v$ is the zenith angle
$i$ and $r$ are the instrument and prism rod heights, respectively, and
$C_{C R}$ is the correction term for curvature and refraction, which is given by:
Equation 4.3 Curvature and refraction correction for trigonometric leveling (after Wolf and Brinker, 1994, p. 110)

$$
C_{C R}=0.0206\left(\frac{D_{S} \sin v}{1000}\right)^{2}[\text { feet }] \quad C_{C R}=0.0675\left(\frac{D_{S} \sin v}{1000}\right)^{2}[\text { meters }]
$$

Note that $C_{C R}$ is always added to the change in elevation computed in Equation 4.2.


## Example computation

Given: Two survey stations (CAS-1 and CAS-4) were occupied with a 1 -second Kern theodolite. The zenith angle was measured from both stations simultaneously in two sets of forward and reverse (face 1, face 2) observations, for a total of 12 measurements. The mean observed zenith angle for the forward and reverse sets are given below, along with the instrument and target heights. The slope distance between CAS-1 and CAS-4 is 2016.615 ift.

| Sets | From | To | Instrument <br> height (ft) | Target <br> height (ft) | Mean zenith angle of all <br> three sets |
| :---: | :---: | :---: | :---: | :---: | :---: |
| $1-3$ | CAS-1 | CAS-4 | 4.89 | 5.21 | $96^{\circ} 37^{\prime} 39.0^{\prime \prime}$ |
| $4-6$ | CAS-4 | CAS-1 | 5.21 | 4.89 | $83^{\circ} 22^{\prime} 41.3^{\prime \prime}$ |

Find: The elevation change from CAS-1 to CAS-4, both with and without correction for curvature and refraction.

Computations:
The correction for curvature and refraction between these stations is essentially the same, since it is based on the horizontal distance, and so any of the observed zenith angles can be used:

$$
C_{C R}=0.0206\left(\frac{D_{S} \sin v}{1000}\right)^{2}=0.0206 \times(=\times \sin (\bar{\square}))^{2}=\square
$$

Now compute each elevation change:

$\Delta H_{1-3}=$ $\qquad$ $\times \cos ($ $\qquad$ ) + $\qquad$ $-$ $\qquad$ $+$ $\qquad$
$\qquad$

## Average height change between stations:

## Solution:

The correction for curvature and refraction between these stations is essentially the same, since it is based on the horizontal distance, and so any of the six zenith angles can be used:

$$
C_{C R}=0.0206 \times\left(\frac{2016.615 \times \sin \left(96.62750^{\circ}\right)}{1000}\right)^{2}=0.0206 \times 2.0031^{2}=\underline{0.083 \mathrm{ft}}
$$

Now compute each elevation change:

$$
\begin{array}{lcc}
\quad \begin{array}{l}
\text { Uncorrected }
\end{array} & \begin{array}{c}
\text { Corrected } \\
\Delta H_{1-3}=2016.615 \times \cos \left(96.62750^{\circ}\right)+4.89-5.21= \\
-233.065 \mathrm{ft} \\
-2.083
\end{array}= & \underline{-232.982 \mathrm{ft}} \\
\Delta H_{4-6}=2016.615 \times \cos \left(83.37814^{\circ}\right)+5.21-4.89=\underline{+232.868 \mathrm{ft}}+0.083= & \underline{+232.951 \mathrm{ft}} \\
\underline{\text { Average height change between stations: }} & \underline{\mathbf{2 3 2 . 9 7} \mathbf{f t}} & \underline{\mathbf{2 3 2 . 9 7} \mathbf{f t}}
\end{array}
$$

The average of uncorrected height changes is the same since effect of curvature and refraction cancels when observations are made in both directions simultaneously.

## Exercise 4.3: Dynamic heights and geopotential numbers

In addition to orthometric heights, $H$ ("elevations"), NGS Datasheets also give dynamic heights, $H^{D}$. A dynamic "height" is actually not a height in the geometric sense of a distance above a reference surface. Rather, it is a geopotential number, $C$, that has been divided (scaled) by a constant value of gravity, which gives $H^{D}$ units of length. Both $C$ and $H^{D}$ represent the gravitational potential energy at a point, and changes in $H^{D}$ are the only "height" differences that give true change in hydraulic head. That is, unconfined water will not flow from one point to another if the water surface at both points has the same $H^{D}$, even though the points will generally not have the same "elevation", $H$ (i.e., $\Delta H^{D} \neq \Delta H$, although the difference is often small).

Equation 4.4 Relationship between dynamic height and geopotential number

$$
\left.H^{D}=\frac{C}{\gamma_{0}} \quad H^{D}=\frac{C}{9.806199} \text { [meters }\right] \quad H^{D}=\frac{C}{32.172569} \text { [ift] }
$$

where $C=$ geopotential number (units of $\mathrm{m}^{2} / \mathrm{s}^{2}$ or $\mathrm{ft}^{2} / \mathrm{s}^{2}$ )
$\gamma_{0}=9.806199 \mathrm{~m} / \mathrm{s}^{2}=$ normal gravity on the GRS 80 ellipsoid at $45^{\circ}$ latitude (given on NGS Datasheets as 980.6199 gals, where $1 \mathrm{~m} / \mathrm{s}^{2}=100$ gals)

Both the dynamic and orthometric heights shown on NGS Datasheets were originally computed from the same set of adjusted geopotential numbers. The relationship between these two types of heights is given below.

Equation 4.5 Relationship between NAVD 88 dynamic and Helmert orthometric heights

$$
H^{D}=\frac{H}{\gamma_{0}} \bar{g}=\frac{H}{\gamma_{0}}\left(g+\frac{H}{K}\right)=\frac{H}{\gamma_{0}}\left(g+\frac{H}{2,358,000}\right)
$$

(modified from Zilkoski et al., 1992)
where $\bar{g}=$ Helmert mean gravity on the plumbline
$g=$ "Observed" (modeled) NAVD 88 surface gravity (given on NGS Datasheets in milligals, where $1 \mathrm{~m} / \mathrm{s}^{2}=100,000 \mathrm{mgals}$ )
$K=2,358,000 \mathrm{~s}^{2}=1 /\left(4.24 \times 10^{-7} \mathrm{~s}^{-2}\right)$ is a constant factor for computing Helmert NAVD 88 mean gravity (assumes constant topographic density of $2670 \mathrm{~kg} / \mathrm{m}^{3}$ )

Equations 4.4 and 4.5 show that orthometric heights can also be computed from geopotential numbers, as $H=C / \bar{g}$.

## Example computation

Given: The NGS Datasheet for NGS station PEND (in Exercise 4.1, and on the next page):
Find: The geopotential number of PEND from both the dynamic and orthometric height (in ift).

## Computations:

Using the published NAVD 88 dynamic height:

$$
\begin{aligned}
& C=\begin{array}{cc}
\gamma_{0} & \times H^{D} \\
C= \\
& \times \frac{\mathrm{m}}{0.3048 \mathrm{~m} / \mathrm{ift}}=
\end{array} \mathrm{ift}^{2} / \mathrm{s}^{2}
\end{aligned}
$$

Using the published NAVD 88 Helmert orthometric height:

$$
\begin{aligned}
& C=\left(g+\frac{H}{K}\right) \times H \\
& C=(\square+\bar{\square}) \times=\frac{\mathrm{m}^{2} / \mathrm{s}^{2}}{(0.3048)^{2} \mathrm{~m}^{2} / \mathrm{ift}^{2}}= \\
& = \\
& \mathrm{ift}^{2} / \mathrm{s}^{2}
\end{aligned}
$$



## Solution:

Using the published NAVD 88 dynamic height:

$$
C=32.172569 \mathrm{ift} / \mathrm{s}^{2} \times \frac{2157.097 \mathrm{~m}}{0.3048 \mathrm{~m} / \mathrm{ift}}=\underline{\mathbf{2 2 7}, \mathbf{6 8 8}} \mathbf{i f t}^{2} / \mathbf{s}^{\mathbf{2}}
$$

Using the published NAVD 88 Helmert orthometric height:

$$
C=\left(9.791254 \mathrm{~m} / \mathrm{s}^{2}+\frac{2160.187 \mathrm{~m}}{2,358,000 \mathrm{~s}^{2}}\right) \times 2160.187 \mathrm{~m}=\frac{21,152.92 \mathrm{~m}^{2} / \mathrm{s}^{2}}{(0.3048)^{2} \mathrm{~m}^{2} / \mathrm{ift}^{2}}=\underline{\mathbf{2 2 7}, \mathbf{6 8 8}} \mathbf{i f t}^{2} / \mathbf{s}^{\mathbf{2}}
$$

## Exercise 4.4: Computing orthometric and dynamic heights from leveling

Leveling, by itself, does not yield true change in orthometric or dynamic heights. But when leveling is combined with surface gravity, the change in geopotential numbers can be computed. If the geopotential number is known for at least one point in a leveling network, then it can be computed at all points in the network. The geopotential numbers can then be converted to orthometric and dynamic heights using the relationships from the previous section, where orthometric height is $H=C / \bar{g}$, and dynamic height is $H^{D}=C / \gamma_{0}$.

Equation 4.6 Determining change in geopotential from leveled height differences

$$
C_{B} \approx C_{A}+\left(\frac{g_{A}+g_{B}}{2}\right) \Delta n_{A B}
$$

where $g_{A}$ and $g_{B}=$ surface gravity at adjacent stations $A$ and $B\left(\mathrm{in} \mathrm{m} / \mathrm{s}^{2}\right.$ or $\left.\mathrm{ft} / \mathrm{s}^{2}\right)$
$\Delta n_{A B}=$ leveled height difference from station $A$ and $B$ (in same linear units as gravity)
Alternatively, leveled height differences can be converted to orthometric heights and dynamic heights by adding an orthometric correction ( $O C$ ) or dynamic correction ( $D C$ ) to observed leveled height differences between adjacent stations.

Equation 4.7 The NAVD 88 Helmert orthometric correction for leveled height differences

$$
O C_{A B} \approx \frac{\left[K\left(g_{A}-g_{B}\right)-2 \Delta n_{A B}\right]\left[2 H_{A}+\Delta n_{A B}\right]}{2\left(K g_{B}+H_{A}+\Delta n_{A B}\right)} \text { (modified from Hwang and Hsiao, 2003) }
$$

where all variables are as defined previously, and the orthometric correction is added to the observed leveled height difference, i.e., $H_{B} \approx H_{A}+\Delta n_{A B}+O C_{A B}$.

Equation 4.8 The dynamic correction for leveled height differences

$$
D C_{A B} \approx\left(\frac{g_{A}+g_{B}}{2 \gamma_{0}}-1\right) \Delta n_{A B} \quad \text { (modified from Hofmann-Wellenhof and Moritz, 2005) }
$$

where all variables are as defined previously, and the dynamic correction is added to the observed leveled height difference, i.e., $H_{B}^{D} \approx H_{A}^{D}+\Delta n_{A B}+D C_{A B}$.
"Approximately equal" symbols were used for equations $4.6-4.8$ because the surface gravity varies continuously along the leveling route. These equations will be exactly true only when the gravity varies linearly between stations. For best results they should be applied to every turning point on a leveling route. However, in most cases, Equation 4.7 (orthometric corrections) should work well for stations less than about 2 km apart. Equations 4.6 and 4.8 (geopotential numbers and dynamic corrections) are more sensitive to variation in surface gravity, and may not give good results even for stations less than 2 km apart, especially in mountainous areas.

## Example computation

Given: A leveled height difference of +50.387 ft measured from NGS stations M 504 (PID FQ0543) to L 504 (PID FQ0544). The following data apply to these stations:

|  | M 504 $($ station $\boldsymbol{A})$ | L 504 $($ station B) |
| :--- | :---: | :---: |
| Orthometric height | 6104.396 ift | $?$ |
| Dynamic height | 6095.991 ift | $?$ |
| Surface gravity | $32.125673 \mathrm{ift} / \mathrm{s}^{2}$ | $32.125305 \mathrm{ift} / \mathrm{s}^{2}$ |

Find: The orthometric and dynamic heights of L 504 (in ift). The stations are 6450 ft apart.
Computations: The stations are (slightly) less than about 2 km apart, so using gravity values only at the stations themselves should be adequate (rather than at every leveling turning point).

Alternative 1: Solve using geopotential numbers.

$$
\begin{aligned}
& C_{B}=C_{A}+\left(\frac{g_{A}+g_{B}}{2}\right) \Delta n_{A B}=\gamma_{0} H_{A}^{D}+\left(\frac{g_{A}+g_{B}}{2}\right) \Delta n_{A B} \\
& C_{B}=+\quad+\left(\frac{+}{2}\right) \times
\end{aligned}
$$

$\qquad$

Orthometric height:

$$
C_{B}=\ldots \mathrm{ift}^{2} / \mathrm{s}^{2}
$$

$$
H_{B}=\frac{C_{B}}{\bar{g}_{B}}=\frac{C_{B}}{g_{B}+\frac{H_{A}+\Delta n_{A B}}{K}}=
$$

$\qquad$


Dynamic height:

$$
H_{B}^{D}=\frac{C_{B}}{\gamma_{0}}=\square=
$$

Alternative 2: Solve using dynamic and orthometric corrections.

$$
\begin{aligned}
& \left.O C_{A B}=\frac{[\quad \times(\ldots)-2 \times \ldots \times[2 \times \ldots}{2 \times(\ldots+\ldots}+\ldots\right) \\
& O C_{A B}= \\
& D C_{A B}=\left(\frac{g_{A}+g_{B}}{2 \times \gamma_{0}}-1\right) \times \Delta n_{A B} \\
& D C_{A B}=(\overline{2 \times \ldots}-1) \times
\end{aligned}
$$

## Orthometric height:

$$
H_{B} \approx H_{A}+\Delta n_{A B}+O C_{A B}=
$$

$\qquad$ $+$ $\qquad$ $+$ $\qquad$ $=$ $\qquad$ ift

## Dynamic height:

$$
H_{B}^{D} \approx H_{A}^{D}+\Delta n_{A B}+D C_{A B}=
$$

$\qquad$ $+$ $\qquad$
$\qquad$ $=$ $\qquad$ ift

Solution:
Alternative 1: Solve using geopotential numbers.

$$
\begin{aligned}
C_{B}=32.172569 \times 6095.991+\left(\frac{32.125673+32.125303}{2}\right) \times 50.387 \\
C_{B}=196,123.7+32.125489 \times 50.387=\underline{197,742.4} \mathrm{ift}^{2} / \mathrm{s}^{2}
\end{aligned}
$$

Orthometric height: $\quad H_{B}=\frac{C_{B}}{\bar{g}_{B}}=\frac{197,742.4}{32.125305+\frac{6104.396+50.387}{2,358,000}}=\underline{\mathbf{6 1 5 4 . 8 4 7} \mathbf{~ i f t}}$
Dynamic height: $\quad H_{B}^{D}=\frac{C_{B}}{\gamma_{0}}=\frac{197,742.4}{32.172569}=$
6146.304 ift

Alternative 2: Solve using orthometric and dynamic corrections.

$$
\begin{array}{r}
O C_{A B}=\frac{[2,358,000 \times(32.125673-32.125305)-2 \times 50.387] \times[2 \times 6104.396+50.387]}{2 \times(2,358,000 \times 32.125305+6104.396+50.387)} \\
O C_{A B}=\frac{[766.970] \times[12,259.179]}{151,515,248}=+\frac{+0.062 \mathrm{ft}}{} \\
D C_{A B}=\left(\frac{32.125673+32.125305}{2 \times 32.172569}-1\right) \times 50.387=(-0.001463) \times 50.387=\quad-0.074 \mathrm{ft}
\end{array}
$$

Orthometric height:

$$
H_{B}=H_{A}+\Delta n_{A B}+O C_{A B}=6104.396+50.387+0.062=\underline{\mathbf{6 1 5 4 . 8 4 5}} \mathbf{i f t}
$$

Dynamic height:

$$
H_{B}^{D}=H_{A}^{D}+\Delta n_{A B}+D C_{A B}=6095.991+50.387+(-0.074)=\underline{\mathbf{6 1 4 6} .304} \mathbf{i f t}
$$

Check: The NGS Datasheet for station L 504 gives:

$$
H_{B}=1875.997 \mathrm{~m}=\underline{6154.846 \mathrm{ift}} \quad \text { and } \quad H_{B}^{D}=1873.393 \mathrm{~m}=\underline{6146.302 \mathrm{ift}}
$$

These results are essentially equal to the NGS Datasheet values, to within the displayed precisions ( $\pm 0.0005 \mathrm{~m}= \pm 0.0016 \mathrm{ft}$ ). However, part of the difference is likely also due to nonlinear variation in gravity between the stations, which are $6450 \mathrm{ft}(1.95 \mathrm{~km})$ apart

Note that $\Delta H=50.449 \mathrm{ft}$ does not equal $\Delta H^{D}=50.311 \mathrm{ft}$, and that only $\Delta H^{D}$ gives true change in hydraulic head (even though it is not really a change in "height", at least in the geometric sense).

## Section 5

## DOCUMENTATION AND ACCURACY REPORTING

## Is it in the right place? By how much? How do you know?

## Examples of documentation and accuracy reporting errors

Table 5.1 Examples of various positioning error sources and their magnitudes due to documentation and accuracy reporting problems (abbreviations and technical terms are defined in the Glossary).

| Documentation error examples | Problem |
| :--- | :---: |
| Documenting geodetic datum as "WGS-84" when data <br> actually referenced to NAD 83 | Perpetuates confusion about <br> "equivalence" of WGS-84 and <br> NAD 83 |
| Listing grid coordinates (such as SPCS) as "NAD 83"" | NAD 83 is a geodetic datum, not <br> a grid coordinate system |
| Documenting geodetic datum as "GRS-80" | GRS-80 is a reference ellipsoid, <br> not a datum |
| Documenting vertical datum as "Mean Sea Level" (MSL) | There is no MSL datum in the <br> US (name changed to <br> NGVD 29 in 1976) |
| Using precision as an accuracy estimate with data containing <br> systematic errors (e.g., incorrect reference coordinates) | Accuracy estimate is <br> meaningless |
| Reporting horizontal error using unscaled standard deviation, <br> rather than at the 95\% confidence level (as specified by the <br> FGDC) | Gives error estimates at 39\% <br> confidence level |
| Reporting vertical error using unscaled standard deviation, <br> rather than at the 95\% confidence level (as specified by the <br> FGDC) | Gives error estimates at 68\% <br> confidence level |
| Using radial and circular estimates for horizontal error rather <br> than semi-major axis of horizontal error ellipse | Typically makes errors appear <br> less than actual |
| Using trivial vectors in GPS network adjustments | Varies, but always makes errors <br> appear less than actual |
| Relying on precision computed by baseline processor for a <br> single GPS vector as an indicator of accuracy | Varies, but precision value <br> usually very optimistic and will <br> not reveal systematic errors |

## Exercise 5.1: Computing error circle and ellipse from standard error components

Accuracies are given on the NGS Datasheet as linear values for the north, east, and up components (in centimeters) scaled to the $95 \%$ confidence level. The north and east components can be converted to a horizontal (circular) accuracy consistent with the approach used by the National Standard for Spatial Data Accuracy (NSSDA) as developed by the Federal Geographic Data Committee (1998, Part 3). Error ellipse axes and rotation can also be computed from the north and east standard error components and horizontal correlation given in the NGS Readjustment Distribution Format (RDF) file.

Equation 5.1 Horizontal (circular) accuracy computed from north and east accuracy components (at the $95 \%$ confidence level per NSSDA)

$$
C E P_{95}=1.2489 \frac{E_{95}^{N}+E_{95}^{E}}{2}=2.4477 \frac{\sigma_{N}+\sigma_{E}}{2}
$$

where $C E P_{95}$ is the estimated Circular Error Probable (horizontal accuracy) at $95 \%$ confidence $E_{95}^{N}$ and $E_{95}^{E}$ are the north and east errors (accuracies), respectively, from the NGS Datasheet (which are given at $95 \%$ confidence)
$\sigma_{N}$ and $\sigma_{E}$ are the north and east standard errors, respectively, from the NGS RDF file
The value 1.2489 is the ratio of the bivariate and univariate scalars for a confidence level of $95 \%$, because the NGS Datasheet gives the north and east accuracies using the univariate scalar at $95 \%$ confidence (see Table 5.2 below for these scalars at this and other confidence levels). Note that CEP is typically computed at the $50 \%$ confidence level.

Equation 5.2 Horizontal covariance computed from correlation (the horizontal correlation is given in the NGS RDF file)

$$
\sigma_{N E}=\rho \sigma_{N} \sigma_{E}
$$

where $\sigma_{N E}$ is the horizontal covariance
$\rho$ is the horizontal correlation

Equation 5.3 Horizontal error ellipse axes computed from standard errors and covariance (at $95 \%$ confidence; standard error values are given in the NGS RDF file)

$$
a, b=2.4477 \sqrt{\frac{1}{2}\left[\sigma_{N}^{2}+\sigma_{E}^{2} \pm \sqrt{\left(\sigma_{N}^{2}-\sigma_{E}^{2}\right)^{2}+4 \sigma_{N E}^{2}}\right]}
$$

where $a$ and $b$ are the error ellipse semi-major and semi-minor axes, scaled to $95 \%$ confidence (note that the " $\pm$ " operator allows computation of both $a$ and $b$ with this one equation, and that $a$ is always greater than $b$ ).

Equation 5.4 Horizontal error ellipse rotation computed from standard errors and covariance (standard error values are given in the NGS RDF file)

$$
\theta=\frac{1}{2} \tan ^{-1}\left(\frac{2 \sigma_{N E}}{\sigma_{E}^{2}-\sigma_{N}^{2}}\right)
$$

where $\theta$ is the rotation angle of the semi-major axis, with respect to the east direction (positive counterclockwise). If $\sigma_{N}>\sigma_{E}$, rotation is with respect to the positive east axis. If $\sigma_{N}<\sigma_{E}$, rotation is with respect to the negative east axis. If $\sigma_{N}=\sigma_{E}$, then $\theta= \pm 45^{\circ}$, where the sign of the rotation is determined by the sign of $\sigma_{N E}$.

Table 5.2 Values used to scale standard errors (accuracies) to various confidence levels. The univariate scalar is used for single error components, such as vertical error. The bivariate scalar is used for dual (two-dimensional) error components, such as horizontal error, and can be used to scale an error ellipse to a desired confidence level. The trivariate scalar is rarely used but is provided here for the sake of completeness. It is for three-dimensional error components and can be used for scaling an error ellipsoid to a desired confidence level. In all cases, these scalars are based on the normal probability distribution of random variables, and the multivariate scalars are for jointly distributed random variables.

| Univariate scalars |  | Bivariate scalars |  | Trivariate scalars |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Scalar, <br> $c_{X}^{1}$ | Confidence <br> level, $X$ | Scalar, <br> $c_{X}^{2}$ | Confidence <br> level, $X$ | Scalar, <br> $c_{X}^{3}$ | Confidence <br> level, $X$ |
| 0.6745 | $50.00 \%$ | 1.0000 | $39.35 \%$ | 1.0000 | $19.87 \%$ |
| 1.0000 | $68.27 \%$ | 1.1774 | $50.00 \%$ | 1.5382 | $50.00 \%$ |
| 1.6449 | $90.00 \%$ | 2.0000 | $86.47 \%$ | 2.0000 | $73.85 \%$ |
| 1.9600 | $95.00 \%$ | 2.1460 | $90.00 \%$ | 2.5003 | $90.00 \%$ |
| 2.0000 | $95.45 \%$ | 2.4477 | $95.00 \%$ | 2.7955 | $95.00 \%$ |
| 2.5758 | $99.00 \%$ | 3.0000 | $98.89 \%$ | 3.0000 | $97.07 \%$ |
| 3.0000 | $99.73 \%$ | 3.0349 | $99.00 \%$ | 3.3682 | $99.00 \%$ |
| 3.2905 | $99.90 \%$ | 3.7169 | $99.90 \%$ | 4.0331 | $99.90 \%$ |

Again, for the sake of completeness, note that the trivariate scalar can be used to scale the estimated Spherical Error Probable (SEP) to a desired confidence level. The estimated SEP at $95 \%$ confidence is computed as

$$
S E P_{95}=2.7955 \frac{\sigma_{N}+\sigma_{E}+\sigma_{U}}{3}=1.4263 \frac{E_{95}^{N}+E_{95}^{E}+E_{95}^{U}}{3}
$$

where $\sigma_{U}$ is the up (ellipsoid height) standard error and $E_{95}^{U}$ is the up error (accuracy) at $95 \%$ confidence as given on the NGS Datasheet. As with CEP, typically SEP is computed at $50 \%$ confidence.

## Example computation

Given: The NGS Datasheet network accuracies and Readjustment Distribution Format (RDF) standard errors (sigmas) for station NN 03 (CZ2412):


Readjustment Distribution Format values for station NN 03 (in centimeters):


Find: The ellipsoid height (up) error, circular error probable (CEP), spherical error probable (SEP), and the horizontal error ellipse axes and rotation angle, all at the $50 \%, 95 \%$, and $99 \%$ confidence levels. Compute all values from either or both the accuracy values on the datasheet or the RDF output, as appropriate, and give the final results in feet.

Computations: The ellipsoid height accuracy (error) is one-dimensional, so the univariate scalars from Table 5.2 should be used to scale the errors to the required confidence levels. This can be done from either the RDF file or the Datasheet.

## Using the RDF sigma value

$$
E_{X}^{U}=c_{X}^{1} \times \sigma_{U}
$$

Using the Datasheet accuracy

$$
E_{X}^{U}=\frac{c_{X}^{1}}{c_{95}^{1}} \times E_{95}^{U}
$$

where $c_{X}^{1}$ is the univariate scalar at the $X \%$ confidence level, and we have:

$$
\begin{aligned}
& E_{50}^{U}=\left\{\begin{array}{c}
0.6745 \times 1.10 \\
o r \\
0.6745 / 1.9600 \times 2.16
\end{array}\right\}=0.74 \mathrm{~cm}=\underline{\mathbf{0 . 0 2 4} \mathbf{f t}(\text { at } \mathbf{5 0 \%} \text { confidence) }} \\
& E_{95}^{U}=\left\{\begin{array}{c}
1.9600 \times 1.10 \\
o r \\
1.9600 / 1.9600 \times 2.16
\end{array}\right\}=2.16 \mathrm{~cm}=\underline{\mathbf{0 . 0 7 1} \mathbf{~ t t}(\text { at 95\% confidence) }} \\
& E_{99}^{U}=\left\{\begin{array}{c}
2.5758 \times 1.10 \\
o r \\
2.5758 / 1.9600 \times 2.16
\end{array}\right\}=2.84 \mathrm{~cm}=\underline{\mathbf{0 . 0 9 3} \mathbf{~ f t ~ ( a t ~ 9 9 \%} \text { confidence) }}
\end{aligned}
$$

The $C E P$ is two-dimensional, so the bivariate scalars from Table 5.2 should be used to scale the errors to the required confidence levels. This can also be done from either the RDF file or the Datasheet.

Using the RDF sigma values

$$
C E P_{X}=c_{X}^{2} \frac{\sigma_{N}+\sigma_{E}}{2} \quad \text { or }
$$

## Using the Datasheet accuracies

$$
C E P_{X}=\frac{c_{X}^{2}}{c_{95}^{1}} \times \frac{E_{95}^{N}+E_{95}^{E}}{2}
$$

where $c_{X}^{2}$ is the bivariate scalar at the $X \%$ confidence level (note that the univariate scalar is used in the denominator for Datasheet accuracies). First we can compute the mean north and east sigma and Datasheet accuracy values as $\left(\sigma_{N}+\sigma_{E}\right) / 2=(0.60+0.61) / 2=\underline{0.605 \mathrm{~cm}}$ and $($ $\left.E_{95}^{N}+E_{95}^{E}\right) / 2=(1.18+1.20) / 2=\underline{1.190 \mathrm{~cm}}$. The $C E P$ for each case is then:

$$
\begin{aligned}
& C E P_{50}=\left\{\begin{array}{c}
1.1774 \times 0.605 \\
\text { or } \\
1.1774 / 1.9600 \times 1.19
\end{array}\right\}=0.71 \mathrm{~cm}=\underline{\mathbf{0 . 0 2 3} \mathbf{~ f t ~ ( a t ~} \mathbf{5 0 \%} \mathbf{c o n f i d e n c e})} \\
& C E P_{95}=\left\{\begin{array}{c}
2.4477 \times 0.605 \\
\text { or } \\
2.4477 / 1.9600 \times 1.19
\end{array}\right\}=1.49 \mathrm{~cm}=\underline{\mathbf{0 . 0 4 9} \mathbf{~ f t ~ ( a t ~ 9 5 \%} \text { confidence) }} \\
& C E P_{99}=\left\{\begin{array}{c}
3.0349 \times 0.605 \\
\text { or } \\
3.0349 / 1.9600 \times 1.19
\end{array}\right\}=1.84 \mathrm{~cm}=\underline{\mathbf{0 . 0 6 0} \mathbf{~ f t ~ ( a t ~ 9 9 \%} \text { confidence) }}
\end{aligned}
$$

The three-dimensional $S E P$ is computed in a similar manner with the trivariate scalars:

Using the RDF sigma values

$$
S E P_{X}=c_{X}^{3} \frac{\sigma_{N}+\sigma_{E}+\sigma_{U}}{3} \quad \text { or }
$$

## Using the Datasheet accuracies

$$
S E P_{X}=\frac{c_{X}^{3}}{c_{95}^{1}} \times \frac{E_{95}^{N}+E_{95}^{E}+E_{95}^{U}}{3}
$$

where $c_{X}^{3}$ is the trivariate scalar at the $X \%$ confidence level. The mean north, east, and up sigma and Datasheet accuracy values are

$$
\begin{aligned}
& \left(\sigma_{N}+\sigma_{E}+\sigma_{U}\right) / 3=(0.60+0.61+1.10) / 3=\underline{0.770 \mathrm{~cm}} \\
& \left(E_{95}^{N}+E_{95}^{E}+E_{95}^{U}\right) / 3=(1.18+1.20+2.16) / 3=\underline{1.51 \mathrm{~cm}}
\end{aligned}
$$

and we have

$$
\begin{aligned}
& S E P_{50}=\left\{\begin{array}{c}
1.5382 \times 0.770 \\
\text { or } \\
1.5382 / 1.9600 \times 1.51
\end{array}\right\}=1.19 \mathrm{~cm}=\underline{\mathbf{0 . 0 3 9} \mathbf{f t}(\text { at } \mathbf{5 0 \%} \mathbf{~ c o n f i d e n c e )}} \\
& S E P_{95}=\left\{\begin{array}{c}
2.7955 \times 0.770 \\
\text { or } \\
2.7955 / 1.9600 \times 1.51
\end{array}\right\}=2.16 \mathrm{~cm}=\underline{\mathbf{0 . 0 7 1} \mathbf{~ f t ~ ( a t ~ 9 5 \% ~ c o n f i d e n c e ) ~}} \\
& S E P_{99}=\left\{\begin{array}{c}
3.3682 \times 0.770 \\
\text { or } \\
3.3682 / 1.9600 \times 1.51
\end{array}\right\}=2.60 \mathrm{~cm}=\underline{\mathbf{0 . 0 8 5} \mathbf{~ f t ~ ( a t ~ 9 9 \%} \text { confidence) }}
\end{aligned}
$$

The horizontal error ellipse must be computed from the RDF values, because the horizontal correlation is not given on the datasheet. We can compute the horizontal covariance from the RDF correlation value:

$$
\sigma_{N E}=\rho \sigma_{N} \sigma_{E}=-0.18761247 \times 0.60 \mathrm{~cm} \times 0.61 \mathrm{~cm}=-0.06867 \mathrm{~cm}^{2} .
$$

The standard error ellipse axes can now be computed using Equation 5.3 (with the $c_{X}^{2}$ value set to one). Note that there is a " $\pm$ " symbol in the equation - $a$ is computed for the case where " $\pm$ " is " + ", and $b$ is computed for the case where " $\pm$ " is " - ":

$$
\begin{aligned}
a, b & =\sqrt{\frac{1}{2}\left[\sigma_{N}^{2}+\sigma_{E}^{2} \pm \sqrt{\left(\sigma_{N}^{2}-\sigma_{E}^{2}\right)^{2}+4 \sigma_{N E}^{2}}\right]} \\
& =\sqrt{\frac{1}{2}\left[0.60^{2}+0.61^{2} \pm \sqrt{\left(0.60^{2}-0.61^{2}\right)^{2}+4 \times(-0.06867)^{2}}\right]}=\left\{\begin{array}{l}
\underline{a=0.66 \mathrm{~cm}=0.022 \mathrm{ft}} \\
\\
\underline{b=0.55 \mathrm{~cm}=0.018 \mathrm{ft}}
\end{array}\right.
\end{aligned}
$$

Since $c_{X}^{2}=1.0000$ for the previous computations, the $a$ and $b$ dimensions are for the standard error ellipse, which has a confidence level of $39.35 \%$ (as shown in Table 5.2). This can be scaled to the required confidence levels using the appropriate bivariate scalars, as follows:

$$
\begin{aligned}
& a_{50}, b_{50}=1.1774\left\{\begin{array}{l}
\times 0.66 \mathrm{~cm}=0.78 \mathrm{~cm}=\underline{\mathbf{0 . 0 2 5} \mathbf{f t}} \\
\times 0.55 \mathrm{~cm}=0.64 \mathrm{~cm}=\underline{\mathbf{0 . 0 2 1} \mathbf{f t}}
\end{array}\right\} \underline{(\text { at } \mathbf{5 0 \%} \text { confidence) }} \\
& a_{95}, b_{95}=2.4477\left\{\begin{array}{l}
\times 0.66 \mathrm{~cm}=1.61 \mathrm{~cm}=\underline{\mathbf{0 . 0 5 3} \mathbf{f t}} \\
\times 0.55 \mathrm{~cm}=1.33 \mathrm{~cm}=\underline{\mathbf{0 . 0 4 4} \mathbf{f t}}
\end{array}\right\} \underline{(\text { at } 95 \% \text { confidence })} \\
& a_{99}, b_{99}=3.0349\left\{\begin{array}{l}
\times 0.66 \mathrm{~cm}=2.00 \mathrm{~cm}=\underline{\mathbf{0 . 0 6 6} \mathbf{f t}} \\
\times 0.55 \mathrm{~cm}=1.65 \mathrm{~cm}=\underline{\mathbf{0 . 0 5 4} \mathbf{f t}}
\end{array}\right\} \underline{(\text { at 99\% confidence })}
\end{aligned}
$$

The error ellipse rotation is computed using Equation 5.4, as follows:

$$
\theta=\frac{1}{2} \tan ^{-1}\left(\frac{2 \sigma_{N E}}{\sigma_{E}^{2}-\sigma_{N}^{2}}\right)=\frac{1}{2} \tan ^{-1}\left(\frac{2 \times(-0.06867)}{0.61^{2}-0.60^{2}}\right)=\underline{-0.74146 \text { radian }}
$$

Converting this to degrees gives:

$$
\theta=\frac{180^{\circ}}{\pi} \times-0.74146=\underline{\mathbf{- 4 2 . 4 8 2}}
$$

The convention is that right-handed (counterclockwise) rotation is positive. So for this case, the error ellipse is rotated $42.482^{\circ}$ clockwise from cardinal directions.

Note the apparent discrepancy in "horizontal" accuracy between what is given on the Datasheet and what was computed for CEP and the error ellipse axes at the $95 \%$ confidence level. The mean value from the Datasheet 0.86 cm , whereas $C E P_{95}=1.08 \mathrm{~cm}$ (which is the same as the mean of the error ellipse axes at $95 \%$ confidence). The reason these values differ is that the Datasheet values are for each individual north and east component, and so to scale these onedimensional values to $95 \%$ confidence level requires a univariate scalar of 1.9600 . The CEP and error ellipse represent the two-dimensional horizontal accuracies, which requires using the bivariate scalar of 2.4477 to scale these values to $95 \%$ confidence. Because of this, if you want to characterize the accuracies on the Datasheet as horizontal (rather than as individual cardinal components), then the Datasheet values should be scaled by the ratio $2.4477 / 1.9600=1.2489$, as shown in Equation 5.1 and was done in the $C E P_{95}$ computation in this example.

## Surveying \& mapping spatial data requirements \& recommendations

These should be explicitly specified in surveying and mapping projects

## 1. Completely define the coordinate system

a. Linear unit (e.g., International foot, U.S. Survey foot, meter)
i. Use same linear unit for horizontal and vertical coordinates
b. Geodetic datum (recommend North American Datum of 1983)
i. Should include datum "tag" (date), e.g., 1986, 1992 (HARN), 2002.0 (CORS)
ii. WGS 84, ITRF, and NAD 27 are NOT recommended
c. Vertical datum (e.g., North American Vertical Datum of 1988)
i. If GPS used for elevations, recommend using a modern geoid model (e.g., GEOID03)
ii. Recommend using NAVD 88 rather than NGVD 29 when possible
d. Map projection type and parameters (e.g., Transverse Mercator, Lambert Conformal Conic)
i. Special attention required for low-distortion grid (a.k.a. "ground") coordinate systems

1) Avoid scaling of existing coordinate systems (e.g., "modified" State Plane)

## 2. Require direct referencing of the NSRS (National Spatial Reference System)

a. Ties to published control strongly recommended (e.g., National Geodetic Survey control)
i. Relevant component of control must have greater accuracy than positioning method used

1) E.g., B-order (or better) stations for GPS control, $2^{\text {nd }}$ order (or better) for vertical control
b. NGS Continuously Operating Reference Stations (CORS) can be used to reference the NSRS
i. Free Internet GPS post-processing service: OPUS (Online Positioning User Service)
3. Specify accuracy requirements (not precision)
a. Use objective, defensible, and robust methods (published ones are recommended)
i. Mapping and surveying: National Standard for Spatial Data Accuracy (NSSDA)
1) Require occupations ("check shots") of known high-quality control stations
ii. Surveys performed for establishing control or determining property boundaries:
2) Appropriately constrained and over-determined least-squares adjusted control network
3) Beware of "cheating" (e.g., using "trivial" GPS vectors in network adjustment)

## 4. Documentation is essential (metadata!)

a. Require a report detailing methods, procedures, and results for developing final deliverables
i. This must include any and all post-survey coordinate transformations

1) E.g., published datum transformations, computed correction surfaces, "rubber sheeting"
b. Documentation should be complete enough that someone else can reproduce the product
c. For GIS data, recommend that accuracy and coordinate system information be included as feature attributes (not just as separate, easy-to-lose and easy-to-ignore metadata files)

## Example of surveying and mapping documentation (metadata)

Basis of Bearings and Coordinates
Linear unit: International foot (ift)
Geodetic datum: North American Datum of 1983 (2007)
Vertical datum: North American Vertical Datum of 1988 (see below)

## System: Arizona LDP

## Zone: Gila Valley

## Projection: Transverse Mercator

Latitude of grid origin: $32^{\circ} 20^{\prime} 00^{\prime \prime} \mathrm{N}$
Longitude of central meridian: $109^{\circ} 48^{\prime} 00$ " W
Northing at grid origin: 0.000 ift
Easting at central meridian: 200,000.000 ift
Scale factor on central meridian: 1.00014 (exact)
All distances and bearings shown hereon are projected (grid) values based on the preceding projection definition. The projection was defined to minimize the difference between projected (grid) distances and horizontal ("ground") distances at the topographic surface within the design area of this coordinate system.
The basis of bearings is geodetic north. Note that the grid bearings shown hereon (or implied by grid coordinates) do not equal geodetic bearings due to meridian convergence.
Orthometric heights (elevations) were transferred to the site from NGS control station "P 439" (PID CY0725) using GPS with NGS geoid model "GEOID09" referenced to the current published NAVD 88 height of this station ( 889.460 m ).
The survey was conducted using GPS referenced to the National Spatial Reference System. A partial list of point coordinates is given below (additional coordinates are available upon request). Local network accuracy estimates are given at the 95\% confidence level and are based on an appropriately constrained least-squares adjustment of over-determined and statistically independent observations.

Point \#1 "SAFFORD BASE ARP", permanent GPS base (off site)

| Latitude $=32^{\circ} 48^{\prime} 07.31561^{\prime \prime} \mathrm{N}$ | Northing $=170,563.997 \mathrm{ift}$ | Estimated accuracy |
| :--- | :--- | :--- |
| Longitude $=109^{\circ} 42^{\prime} 42.84664$ " W | Easting $=227,075.294 \mathrm{ift}$ | Horizontal = Fixed |
| Ellipsoidal height $=2945.423$ ift | Elevation $=3033.826 \mathrm{ift}$ | Vertical = Fixed |

Point \#1002, 1/2" rebar with aluminum cap, derived coordinates (on site)
Latitude $=32^{\circ} 50^{\prime} 06.81662^{\prime \prime} \mathrm{N} \quad$ Northing $=182,643.211 \mathrm{ift} \quad$ Estimated accuracy
Longitude $=109^{\circ} 42^{\prime} 47.90144^{\prime \prime} \mathrm{W} \quad$ Easting $=226,633.861 \mathrm{ift} \quad$ Horizontal $= \pm 0.034 \mathrm{ift}$
Ellipsoidal height $=2822.412$ ift $\quad$ Elevation $=2910.734 \mathrm{ift} \quad$ Vertical $= \pm 0.056 \mathrm{ift}$
Point \#1006, 1/2" rebar with plastic cap, derived coordinates (on site)
Latitude $=32^{\circ} 50^{\prime} 16.89645^{\prime \prime} \mathrm{N} \quad$ Northing $=183,662.115 \mathrm{ift}$
Longitude $=109^{\circ} 42^{\prime} 47.93756^{\prime \prime} \mathrm{W}$
Ellipsoidal height $=2815.734 \mathrm{ift}$

Easting $=226,629.942 \mathrm{ift}$
Elevation $=2904.040 \mathrm{ift}$

Estimated accuracy
Horizontal $= \pm 0.047 \mathrm{ift}$
Vertical $= \pm 0.068$ ift

## GLOSSARY

Below is a list of the abbreviations and terms used in this workbook. In the interest of brevity, the definitions are highly general and simplified. Please note also that this list gives only a portion of the terms and abbreviations frequently encountered in GPS positioning and geodesy. Terms in italics within the definitions are also defined in this list. Cited references are listed at the end of the workbook.

Autonomous position. A $G P S$ position obtained with a single receiver using only the ranging capability of the GPS code (i.e., with no differential correction).

Cartesian coordinates. Coordinates based on a system of two or three mutually perpendicular axes. Map projection and ECEF coordinates are examples two- and three-dimensional Cartesian coordinates, respectively.

Confidence interval or level. A computed probability that the "true" value will fall within a specified region (e.g., $95 \%$ confidence level). Applies only to randomly distributed errors.
CORS (Continuously Operating Reference Stations). A nation-wide system of permanently mounted $G P S$ antennas and receivers that collect GPS data continuously. The CORS network is extremely accurate and constitutes the primary survey control for the US. CORS data can be used to correct GPS survey and mapping results, and the data are freely available over the Internet.

Datum transformation. Mathematical method for converting one geodetic or vertical datum to another (there are several types, and they vary widely in accuracy).

Differential correction. A method for removing much of the error in an autonomous GPS position. Typically requires at least two simultaneously operating GPS receivers, with one of the two at a location of known geodetic coordinates.

ECEF (Earth-Centered, Earth-Fixed). Refers to a global three-dimensional (X, Y, Z) Cartesian coordinate system with its origin at the Earth's center of mass, and "fixed" so that it rotates with the solid Earth. The Z-axis corresponds to the Earth's conventional spin axis, and the X- and Y-axes lie in the equatorial plane. Widely used for geodetic and GPS computations.

Ellipsoid. A simple mathematical model of the Earth corresponding to mean sea level (the geoid) and used as part of a geodetic datum definition. Constructed by rotating an ellipse about its semi-minor axis. Also referred to as a "spheroid".

Ellipsoid height. Straight-line height above and perpendicular to the ellipsoid. This is the type of height determined by $G P S$, and it does not equal elevation. Can be converted to orthometric heights ("elevations") using a geoid model.

Ellipsoid normal. A line perpendicular to the reference ellipsoid along which ellipsoid heights are measured.

FBN (Federal Base Network). Nationwide network of GPS control stations observed and adjusted by the NGS. A nation-wide readjustment of the FBN is scheduled for 2007.
FGDC (Federal Geographic Data Committee). Develops and promulgates information on spatial data formats, accuracy, specifications, and standards. Widely referenced by other organizations. Includes the Federal Geodetic Control Subcommittee (FGCS) and the NSSDA.

Geodetic datum. Reference frame for computing geodetic coordinates (latitude, longitude, and ellipsoid height) of a point. A datum always refers to a particular ellipsoid and a specific adjustment (e.g. the 1992 adjustment of $N A D 83$ for the Arizona $H A R N$ ).

Geographic "projection". This is not a true map projection in the sense that it does not transform geodetic coordinates (latitude and longitude) into linear units. However, it is a projection in the sense that it represents geodetic coordinates on a regular flat grid, such that the difference in angular units (e.g., decimal degrees) is equal in all directions. Because of meridian convergence, this results in an extremely distorted coordinate system, especially at high latitudes, and the distortion varies greatly with direction.

Geoid. Surface of constant gravitational equipotential (a level surface) that best corresponds to global mean sea level. Often used as a reference surface for vertical datums.

GPS (Global Positioning System). A constellation of satellites used for navigation, mapping, surveying, and timing. Microwave signals transmitted by the satellites are observed by GPS receivers to determine a three-dimensional position. Accuracy varies greatly depending on the type of receiver and methods used.

Grid distance. The horizontal distance between two points on a flat plane. This is the type of distance obtained from map projections.

Ground distance. The horizontal distance between two points as measured on the curved Earth surface.
GRS-80 (Geodetic Reference System of 1980). The reference ellipsoid currently used for many geodetic datums throughout the world, including NAD 83 and ITRF.

HARN (High Accuracy Reference Network). Network of GPS stations adjusted by the NGS on a state-by-state basis. The Arizona HARN was adjusted in 1992. In some states it is referred to as a High Precision GPS (or Geodetic) Network (HPGN).

International Foot. Linear unit adopted by the US in 1959, and defined such that one foot equals exactly 0.3048 meter. Shorter than the US Survey Foot by 2 parts per million (ppm).

ITRF (International Terrestrial Reference Frame). Global geodetic reference system that takes into account plate tectonics (continental drift) and is used mainly in scientific studies. A new ITRF "epoch" is computed periodically and is referenced to a specific time (e.g., ITRF 2000 1997.0). Each epoch is a realization of the International Terrestrial Reference System (ITRS). See Soler (2007), and Soler and Snay (2004) for information on its relationship to NAD 83 and WGS 84.

Local geodetic horizon. A "northing", "easting", and "up" planar coordinate system defined at a point such that the northing-easting plane is perpendicular to the ellipsoid normal, north corresponds to true geodetic north, and "up" is in the direction of the ellipsoid normal at that point.

Map projection. A functional (one-to-one) mathematical relationship between geodetic coordinates (latitude, longitude) on the curved ellipsoid surface, and grid coordinates (northings, eastings) on a planar (flat) map surface. All projections are distorted, in that the relationship between projected coordinates differs from that between their respective geodetic coordinates. See Snyder (1987) for details.

NAD 27 (North American Datum of 1927). Geodetic datum of the US prior to NAD 83, and superseded by NAD 83 in 1986. This is the datum of SPCS 27 and UTM 27.

NAD 83 (North American Datum of 1983). Current official geodetic datum of the US. Replaced $N A D$ 27 in 1986, which is the year of the initial NAD 83 adjustment. This is the datum of SPCS 83 and UTM 83. See Schwarz (1986) for details.

NADCON. Datum transformation computer program developed by the NGS for transforming coordinates between NAD 27 and NAD 83, and also between the NAD 831986 adjustment and the various HARN adjustments. See Dewhurst (1990) for details.

NAVD 88 (North American Vertical Datum of 1988). Current official vertical datum of the US. Replaced NGVD 29 in 1991. See Zilkoski et al. (1992) for details.
NDGPS (National Differential GPS). A nation-wide system of "beacons" (permanently mounted GPS receivers and radio transmission equipment) that transmits real-time differential corrections which can be
used by GPS receivers equipped with the appropriate radio receivers. Operated and maintained by the US Coast Guard. See US Coast Guard (2004) for details.
NGS (National Geodetic Survey). Federal agency within the Department of Commerce responsible for defining, maintaining, and promulgating the $N S R S$ within the US and its territories.
NGVD 29 (National Geodetic Vertical Datum of 1929). Previous vertical datum of the US, superseded by NAVD 88 in 1991. Not referenced to the geoid or mean sea level, and not as compatible with GPSderived elevations as NAVD 88. Called "Mean Sea Level" (MSL) datum prior to 1976.

NSRS (National Spatial Reference System). The framework for latitude, longitude, height, scale, gravity, orientation and shoreline throughout the US. Consists of geodetic control point coordinates and sets of models describing relevant geophysical characteristics of the Earth, such as the geoid and surface gravity. Defined, maintained, and promulgated by the NGS (see Doyle, 1994, for details).

NSSDA (National Standard for Spatial Data Accuracy). FGDC methodology for determining the positional accuracy of spatial data (see Federal Geographic Data Committee, 1998).

OPUS (Online Positioning User Service). A free NGS service that computes NSRS and ITRF coordinates with respect to the CORS using raw GPS data submitted via the Internet.

Orthometric correction. A correction applied to leveled height differences which reduces systematic errors due to variation in gravitational potential. See Dennis (2004) for details.

Parts per million (ppm). A method for conveniently expressing small numbers, accomplished by multiplying the number by 1 million (e.g., $0.00001=10 \mathrm{ppm}$ ). Exactly analogous to percent, which is "parts per hundred".

SPCS (State Plane Coordinate System). A system of standardized map projections covering each state with one or more zones such that a specific distortion criterion is met (usually 1:10,000). Projection parameters (including units of length) are independently established by the legislature of each state. Can be referenced to either the NAD 83 or NAD 27 datums (SPCS 83 and SPCS 27, respectively). See Stem (1989) for details.

Triangulation. A method for determining positions from angles measured between points (requires at least one distance to provide scale).

Trilateration. A method for determining positions from measured distances only.
Trivial vector. A GPS vector (computed line connecting two GPS stations) that is not statistically independent from other GPS vectors observed at the same time.

US Survey Foot. Linear unit of the US prior to 1959, and defined such that one foot equals exactly 1200 / 3937 meter. Longer than the International Foot by 2 parts per million (ppm).

UTM (Universal Transverse Mercator). A grid coordinate system based on the Transverse Mercator map projection which divides the Earth (minus the polar regions) into 120 zones in order to keep map scale error within 1:2500. Can be referenced to either the NAD 83 or NAD 27 datums (UTM 83 and UTM 27, respectively). See Hager et al. (1989) for details.

Vertical datum. Reference system for determining "elevations", typically through optical leveling. Modern vertical datums typically use the geoid as a reference surface and allow elevation determination using GPS when combined with a geoid model.

WAAS (Wide Area Augmentation System). A system of geosynchronous satellites and ground GPS reference stations developed and managed by the Federal Aviation Administration and used to provide free real-time differential corrections. See Federal Aviation Administration (2003) for details.

WGS 84 (World Geodetic System of 1984). Reference ellipsoid and geodetic datum of GPS, defined and maintained by the US Department of Defense. Current realizations of WGS 84 are considered identical to ITRF 2000 at the 2 cm level. See National Imagery and Mapping Agency (1997) for details, and Merrigan et al. (2002) for information on the most recent realization.

## SELECTED GPS AND GEODESY REFERENCES

Primary resource: The National Geodetic Survey (http://www.ngs.noaa.gov/)
Some NGS web pages of particular interest
Control station datasheets: http://www.ngs.noaa.gov/cgi-bin/datasheet.prl
The Geodetic Tool Kit: http://www.ngs.noaa.gov/TOOLS/
Online Positioning User Service (OPUS): http://www.ngs.noaa.gov/OPUS/
Continuously Operating Reference Stations (CORS): http://www.ngs.noaa.gov/CORS/
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